

A microseismic study in the western part of the Gulf of Corinth (Greece): implications for large-scale normal faulting mechanisms

A. Rigo,^{1,*} H. Lyon-Caen,¹ R. Armijo,¹ A. Deschamps,² D. Hatzfeld,³
K. Makropoulos,⁴ P. Papadimitriou⁴ and I. Kassaras⁴

¹ *Équipe de Sismotectonique – CNRS URA 195 et 1093 – Institut de Physique du Globe de Paris, 4 Place Jussieu, 75252 Paris Cedex 05, France*

² *Institut de Géodynamique, UMR géosciences Ajur, EP125, CNRS-UNSA, 250 rue A. Einstein, 06560 Valbonne, France*

³ *L.G.I.T-I.R.I.G.M., CNRS URA 733, Université Joseph Fourier, BP 68, 38402 St Martin d'Hères Cedex, France*

⁴ *University of Athens, Department of Geology, Geophysics-Geothermy Division, Panepistimiopolis Ilissia, Athens 157 84, Greece*

Accepted 1996 March 19. Received 1996 March 19; in original form 1995 July 17

SUMMARY

We present the results of a dense seismological experiment in the western part of the Gulf of Corinth (Psathopyrgos–Aigion area), one of the most active rifts in the Aegean region for which we have precise tectonic information. The network included 51 digital stations that operated during July and August 1991, covering a surface of 40×40 km². Among the 5000 recorded events with M_L ranging between 1.0 and 3.0, we precisely located 774 events. We obtained 148 well-constrained focal mechanisms using *P*-wave first motions. Of these, 60 also have mechanisms obtained by combining the *P*-wave first motions with the *S*-wave polarization directions. The observed seismicity is mainly located between 6 and 11 km depth. Most of the fault-plane solutions correspond to E–W-striking normal faulting, in agreement with the geological evidence. Most of the well-determined mechanisms indicate a nodal plane dipping 10–25° due north and a steep south-dipping plane. A similar asymmetry is also seen in the seismicity distribution and in the overall geological structure of the Corinth Rift. We discuss this evidence and the inference of a deep detachment zone, a structure where the major faults seen at the surface appear to root. A large part of the microseismic activity appears to cluster in regions near the junctions of the main faults with the proposed detachment zone. This feature of the microseismicity is interpreted in terms of stress transfer and stress concentration in regions of probable nucleation of future large earthquakes.

Key words: fault-plane solutions, Greece, Gulf of Corinth, normal faulting, microseismicity.

1 INTRODUCTION

One important problem in seismotectonics is to characterize the loading process on large active fault systems and to understand the temporal and spatial relationships between small-scale faulting, microseismicity and large destructive earthquakes. How strains accommodate near the base of the seismogenic layer where large earthquakes nucleate is of particular interest in understanding the mechanisms by which stresses build up. In the case of crustal extension, it is commonly accepted that the normal faults capable of generating large earthquakes usually cut the entire seismogenic layer, but how these faults root into the deeper semi-brittle and ductile layers

remains largely unknown. An asymmetric deep structure is often inferred from seismic profiles for continental areas under extension (Covington 1983; Arabasz & Julander 1986; Logatchev & Zorin 1987; Maier & Eisbacher 1991; Wenzel *et al.* 1991). This implies the existence of crustal detachments or large shear zones dipping less than 30°. Studies of large continental normal-faulting earthquakes, however, show that they generally initiate in the upper brittle crust and the fault-plane solutions do not exhibit low-angle nodal planes (e.g. Jackson & White 1989), with the possible exception of the Woodlark–D'Entrecasteaux earthquake (Abers 1991). In some large earthquakes, secondary rupture on low-angle normal faults has also been suggested (e.g. Eyidogan & Jackson 1985; Bernard & Zollo 1989a). However, the study of the deeper parts of the faults is hampered by the insufficiency of well-defined seismicity.

* Now at: Observatoire Midi-Pyrénées–GRGS–CNRS UMR 5562, 14 Ave. E. Belin, 31400 Toulouse, France.

One possible way to define the fault geometry at depth is to study the deformation in the semi-brittle transition zone by means of its microseismic activity. This can be criticized because fault-plane solutions of microearthquakes do not usually show a simple pattern (e.g. Jackson & White 1989). The main difficulties with this method are (1) to appraise the relationship of the observed microseismic activity (recorded during short periods of time) with the seismic cycle, and (2) to ascribe this activity to the major faults or to the network of fractures surrounding them. A way to overcome these problems is to install very dense networks of portable seismological stations in rapidly deforming areas with high seismicity. However, previous experiments in Greece (Melis *et al.* 1989; Hatzfeld *et al.* 1990) and in the Basin and Range (Jones 1987) had no conclusive results about the normal-faulting geometry at depth, possibly because the deployed networks were not dense enough.

The Gulf of Corinth is one of the most prominent active rifts in the Aegean area, with a history of repeated large earthquakes and with a high level of background seismicity (Papazachos & Papazachos 1989; Ambraseys & Jackson 1990). Large earthquakes occurred in 1861 near the city of Aigion ($M_w = 6.7$) (Papazachos & Papazachos 1989), producing 13 km long surface breaks (Schmidt 1881), and in 1981 near Corinth, with a sequence of three events ($M_s = 6.7, 6.4, 6.4$) (Jackson *et al.* 1982; King, Oppenheimer & Amelung 1985) (Fig. 1). The geological evidence suggests an unusually fast Quaternary slip rate of the order of 1 cm yr^{-1} on the major E–W-striking, north-dipping normal faults at the southern edge of the Corinth

rift, which appears to be mostly an asymmetric structure (Armijo *et al.* 1996).

In 1991, we began a multidisciplinary study, which included the tectonics, the geodesy and the microseismicity of the Gulf of Corinth. Detailed accounts of the tectonic results are given elsewhere (Armijo *et al.* 1996), and a synoptic view of the overall work can be found in Rigo (1994). Here, we present the main results of an experiment in July and August 1991 with a dense network of digital seismological stations installed in the western part of the Gulf. The experiment was designed to cover, at the appropriate scale and with reasonably good hypocentral resolution, an area where the near-surface geometry of the main faults was well constrained. Previous background seismicity proved to be high in the region between the cities of Patras, Aigion and Navpaktos (Galanopoulos 1960, 1961; Melis *et al.* 1989; Hatzfeld *et al.* 1990; Amorese 1993). During our six week experiment we recorded a total of about 5000 events, providing us with a consistent data set from which we obtained accurate earthquake locations and well-constrained fault-plane solutions. We use this data together with the tectonic observations to bring constraints on the geometry of the faults and on the mechanism of normal faulting at depth.

2 SEISMOLOGICAL DATA ANALYSIS

2.1 Network and locations

The seismological network, composed of 51 digital portable stations (Fig. 2; Appendix A), covered an area about

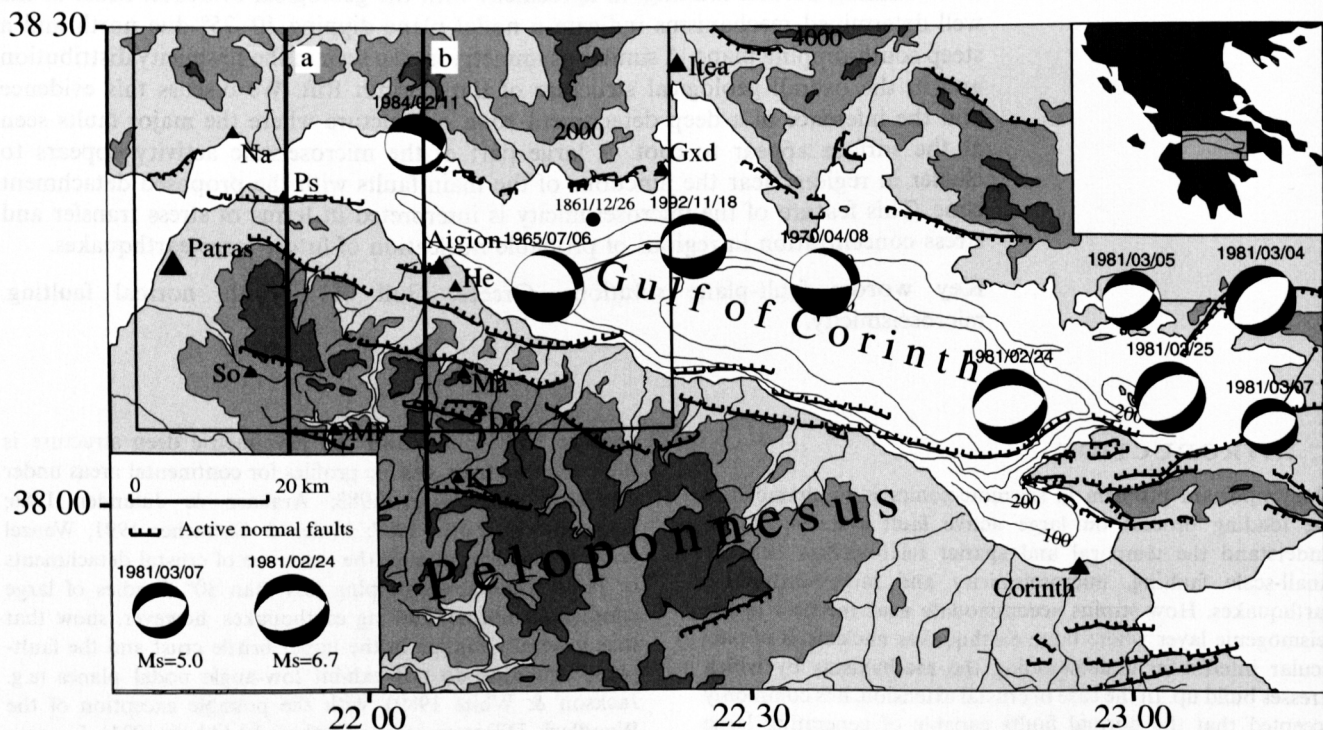


Figure 1. Seismotectonic map of the Gulf of Corinth. Active faults are mapped from field observations and SPOT images (Armijo *et al.* 1996). Elevation contours are in feet. Fault-plane solutions of 1965–92 events with magnitude larger than 5 are from Ambraseys & Jackson (1990), Taymaz, Jackson & McKenzie (1991) and Briole *et al.* (1993). The epicentre of the 1861/12/26 earthquake (black dot) is from Schmidt (1881). a and b indicate locations of sections in Fig. 12. Do: Doumena; Gxd: Galaxidi; He: Helike; Kl: Kalavrita; Ma: Mamousia; MP: Megas Pontias; Na: Navpaktos; Ps: Psathopyrgos; So: Souli. The inset outlines the area studied (Figs 2 and 8).

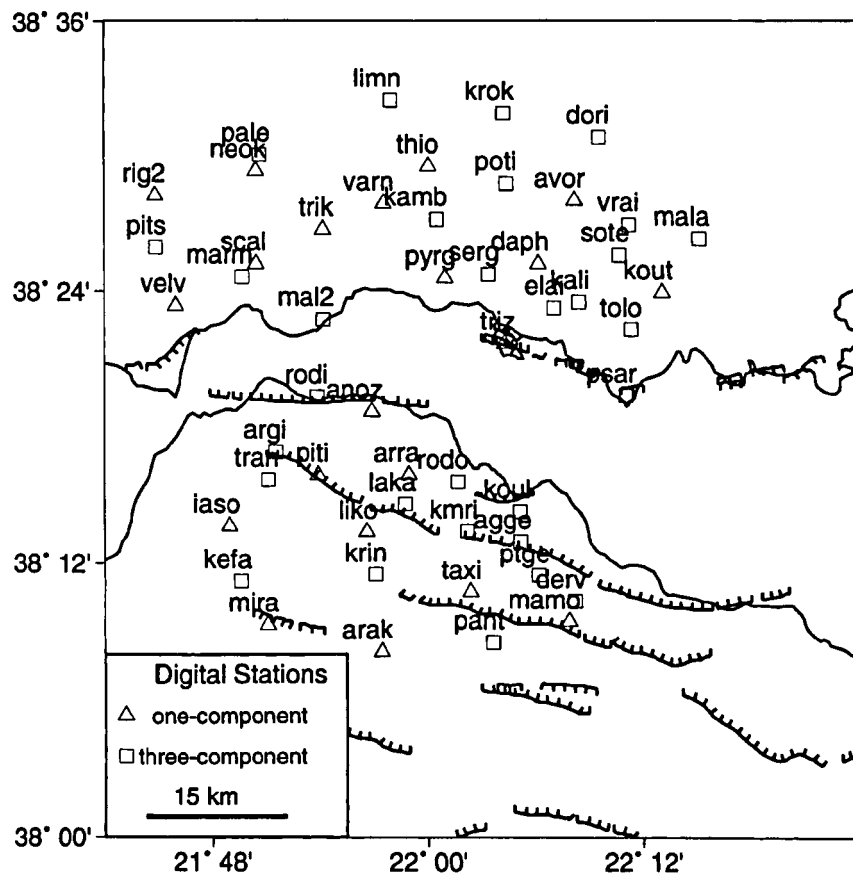


Figure 2. Network in operation from 10/07/1991–26/08/1991. Coordinates of the stations are given in Appendix A.

40 km × 40 km, with an average distance between stations of 3–6 km. The stations were equipped with 1 or 2 Hz seismometers. The 21 one-component stations had a recording frequency of 100 Hz; for the three-component stations, 15 had a recording frequency of 125 Hz, 12 had one of 200 Hz, and 3 had one of 62.5 Hz. Absolute time was given by the DCF time code recorded in parallel with the internal station clock for all three-component stations. For the one-component stations, the internal clock was permanently reset on the DCF time code. Thus the time accuracy is estimated to be better or of the order of the sampling rate. The time-reading precision is estimated to be between 0.005 and 0.06 s for *P* waves, and between 0.01 and 0.1 s for *S* waves, depending on the recording system.

About 5000 events with magnitude M_L between 1 and 3.0 were recorded during the six weeks of the experiment. Here, we analyse a subset of 850 events recorded by at least five stations during the month of August when the network was completely installed. The hypocentral locations were computed using the HYP071 code (Lee & Lahr 1975). As no crustal velocity structure is available for the area, we constructed a mean velocity model from our data set. We estimated the V_p/V_s ratio to be 1.80 ± 0.02 from Wadati diagrams constructed using the three-component records only. This ratio is in agreement with the values ranging between 1.77 and 1.83 obtained in the neighbouring areas (e.g. King *et al.* 1985; Melis *et al.* 1989; Hatzfeld *et al.* 1990; Amorese 1993). We first determined the mean *P* velocity between 0 and 15 km depth by studying the variation of the mean rms travelt ime residual with respect to

this mean velocity (Fig. 3). The mean rms is minimum between 5.4 and 5.7 km s⁻¹. Then, taking a velocity of 5.7 km s⁻¹ as a reference, we computed the number of events for which the rms decreases when using a different velocity. For a value of 5.6 km s⁻¹, 90 per cent of the events have their rms reduced. A similar result was obtained with other reference velocities. We thus used a 5.6 km s⁻¹ mean *P* velocity between 0 and 15 km (model 1—Table 1). Comparable mean velocities were found in the area (e.g. King *et al.* 1985; Melis *et al.* 1989; Amorese 1993). Our data provide no information on the velocity below 15 km depth because of the lack of events at these depths. We thus fixed the velocities there as in the previous studies mentioned above (Table 1). We determined a

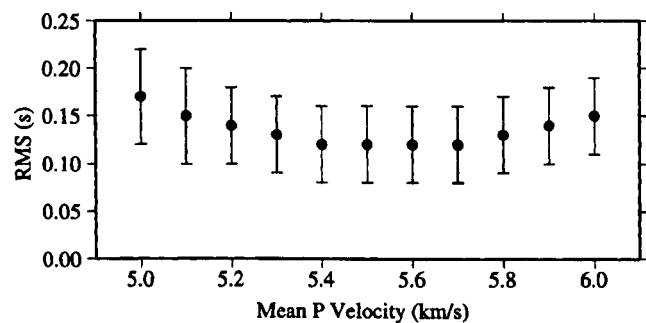


Figure 3. Mean location rms for the 850 selected events versus assumed mean *P*-wave velocities between 0 and 15 km depth. Error bars represent standard deviation.

Table 1. *P*-wave velocity models used for the earthquake locations. $V_p/V_s = 1.80$.

	Model 1	Model 2
Depth (km)	V_p (km/s)	V_p (km/s)
0.0 - 4.0	5.6	4.8
4.0 - 7.2		5.2
7.2 - 8.2		5.8
8.2 - 10.4		6.1
10.4 - 15.0		6.3
15.0 - 30.0	6.5	6.5
> 30.0	7.0	7.0

multilayer model (model 2—Table 1), keeping the 5.6 km s^{-1} mean *P* velocity in the upper 15 km, which appears to be the most appropriate for fault-plane solutions (see detailed discussion in Section 2.2). This model was then used to determine both the event locations and the fault-plane solutions.

Owing to the geometry and the density of the network, event locations are very well constrained. We performed two tests to assess the precision of focal depths. First, we compared results obtained using initial depths of 5 and 20 km with HYPO71. Second, we compared results obtained with velocity models 1 and 2. We rejected events for which the focal depth varied by more than 2 km to select a final set of 774 events with a precision of about $\pm 2 \text{ km}$ for focal depths and $\pm 1 \text{ km}$ for the epicentral positions (Appendix B). The final mean rms travelt ime residual for this final set of events is $0.11 \pm 0.04 \text{ s}$.

2.2 Fault-plane solutions

The fault-plane solutions were determined in two steps. First they were constructed by hand using only the *P* first motions. Then, when it was possible, they were better constrained using *S*-wave polarization directions, following the method developed by Zollo & Bernard (1991). In order to guarantee the quality of the determinations, we selected events with at least 10 clear *P* polarities. This restriction gave us a set of 190 events.

Velocity model 1 did not yield fault-plane solutions for many of these events because the ray's angle of incidence was too small, making it impossible to draw two perpendicular nodal planes. (See, for example, events 398 and 50 in Fig. 4.) The multilayer velocity model 2 was constructed to determine fault-plane solutions for a maximum number of events. As a first step, we proceeded by trial and error, using the 190 selected events, to introduce layers between 0 and 15 km in order to transform some direct rays into refracted rays (e.g. event 50, Fig. 4) which are clearly observed on the records. In the resulting model, the depths of the velocity interfaces are mainly constrained by the depth of the events. Clearly model 2 does not correspond to a unique solution, but it is the best layered model that we could find to reduce polarity inconsistencies in the fault-plane solutions. Tests were performed to assess the stability of the fault-plane solutions with respect to the velocities in the layers and the position of the interfaces. Because of the high density of stations in the network, fault-plane solutions are more sensitive to the location of interfaces than to the velocities in each layer. The 3-D smoothly varying velocity model obtained by tomographic inversion with the same set of data (Le Meur 1994) is consistent with model 2. The interfaces of model 2 correspond to the higher gradients of Le Meur's model. Additionally, focal mechanisms determined with model 1 did not significantly change with model 2 (e.g. event

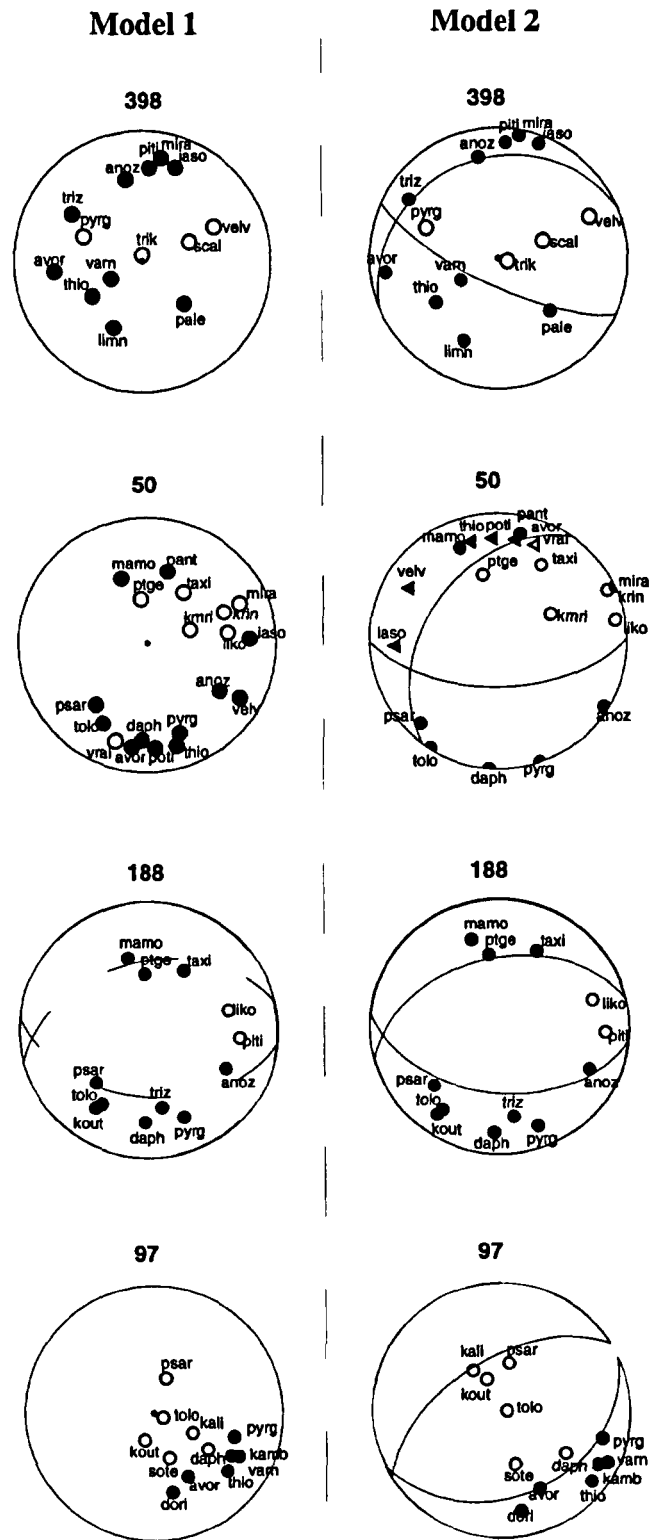


Figure 4. Focal mechanisms of four events (Appendix B) with two different velocity models (models 1 and 2, Table 1). *P* polarities are plotted on the lower equal-area hemisphere of the focal sphere. Dots represent direct rays, triangles refracted rays. Fault planes are drawn when possible. Note that in model 2 the incidence angles are increased and refracted rays are present.

188, Fig. 4). Finally, we obtained a total of 148 focal mechanisms, of which about 30 are not well constrained (e.g. event 97 in Fig. 4).

As a second step to constrain fault-plane solutions better, we combined *P*-wave polarities with *S*-wave horizontal polarization directions to estimate the fault-plane parameters (Zollo & Bernard 1991; de Chabalière *et al.* 1992). For such a study, the *S*-wave polarization must be stable in time and linear. This implies the use of incidence angles smaller than 45° . For larger incidences, free-surface effects can be important, inducing the instability of *S*-wave polarization (Evans 1984; Booth & Crampin 1985; Bernard & Zollo 1989b). To obtain more stable polarizations, the records were bandpass-filtered between 1 and 3 Hz with a zero-phase-shift Butterworth filter. However, few stations show very clear splitting of *S* waves due to crustal anisotropy (Bouin, Tèllez & Bernard 1996) with time delays larger than 0.05 s between the slow and the rapid *S* arrivals. These data were not used in the inversion. Then, each selected three-component record was examined for *S*-polarization stability. In Fig. 5 we present an example of the determination of the *S*-wave polarization direction at station krin for event 447. The initial record (Fig. 5a) is filtered between 1 and 3 Hz (Fig. 5b). We represent the 3-D particle movement vector at a constant time interval on a horizontal stereographic projection for a short temporal window (1 s) for the *P* wave and the *S* wave. We can thus check that the *P*-wave incidence is vertical. If the *S*-wave polarization is stable and linear with time, the points cluster (each point representing the particle vector at a given time on the stereographic projection) around the direction of the *S*-wave polarization (west in the example of Fig. 5). The average values of the azimuths and incidence angles of the particle movement vector at the surface define the *S*-wave mean polarization direction and incidence angle. The nearer to the horizontal the *S* incidence angle is, the better the datum. The accuracy of the polarization direction depends on the size of the particle movement vector cluster, i.e. on the stability of the *S*-wave polarization with time. This accuracy is globally estimated to be $10\text{--}15^\circ$. Thus, for 60 events out of the 190 selected earlier, we were able to add information by including the *S*-wave polarization directions.

We then performed the inversion for each mechanism by searching for the best fault-plane solution corresponding to the observed *P* polarities and *S*-wave polarization directions. Because of the non-linearity of the problem, the inversion is a systematic exploration of the fault-plane parameter space (strike, dip, slip) to find the maximum of a probability density function (Brillinger, Udias & Bolt 1980; Zollo & Bernard 1991). The result of the joint inversion of *P* polarities and *S*-polarization directions for event 491 is shown in Fig. 6. The probability density function is represented by the 66, 90 and 99 per cent levels on the entire fault-plane parameter space. The clusters represent the conjugate fault planes (P1 and P2 in Fig. 6). The estimated values for strike, dip and slip are given by the position of the maximum of the density function for each nodal plane. In the example of Fig. 6, the values are: strike = 211.0° , dip = 70.0° and slip = -150.4° . The uncertainties on these values can be estimated from the area with 66 per cent probability. For a given slip and dip, the uncertainty on the strike is close to 30° ; for a given strike, the uncertainties on slip and dip range between 10° and 30° . Because we have to consider the three fault-plane parameters

Event 447 - Station krin (krini)

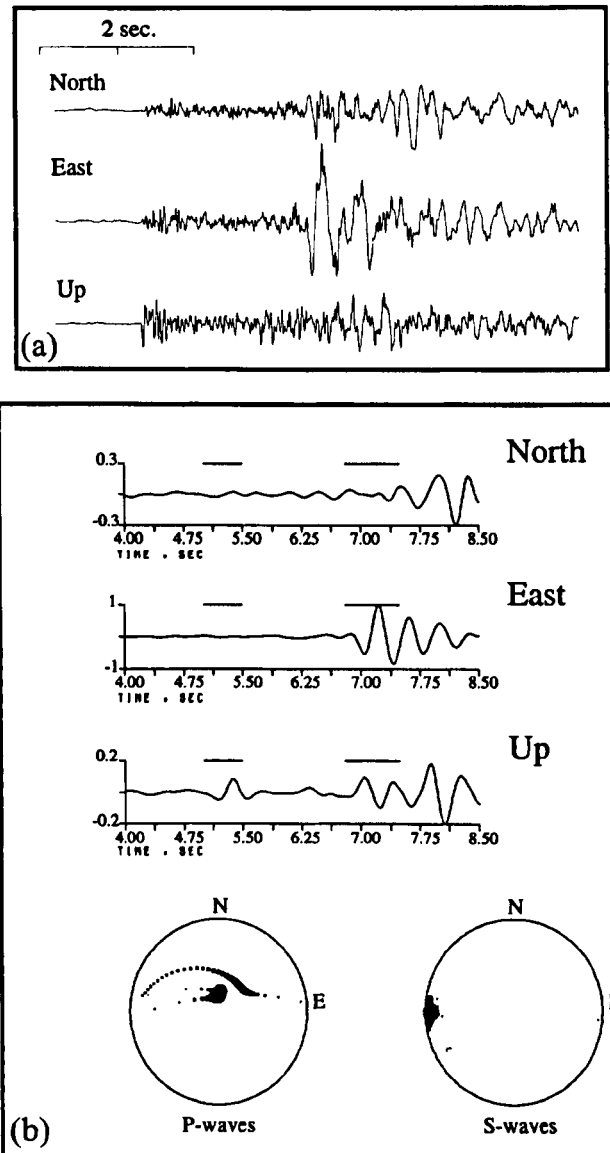


Figure 5. Example of determination of shear-wave direction of polarization. (a) Initial three-component record. (b) Same record filtered between 1 and 3 Hz (top), with particle movement vector on horizontal stereographic projection for *P* waves (bottom-left) and *S* waves (bottom-right). The time window used for *P* and *S* waves is indicated on the filtered records by horizontal lines. The direction and the dip of the *S*-wave polarization obtained here are 90.0° and $7.5^\circ (\pm 3.4^\circ)$, respectively.

together, the overall uncertainty of each of these parameters is about $\pm 15^\circ$.

The four examples in Fig. 7 illustrate the improvement brought by our method. We compare here the mechanisms obtained with the *P* first motions only and the mechanisms from the joint inversion of *P* polarities and *S*-polarization directions. In general, the addition of *S*-polarization directions constrains the mechanisms better. In a few cases, however (e.g. event 447, Fig. 5), the fit of *S*-polarization directions may be slightly worse for the inversion solution than the one obtained

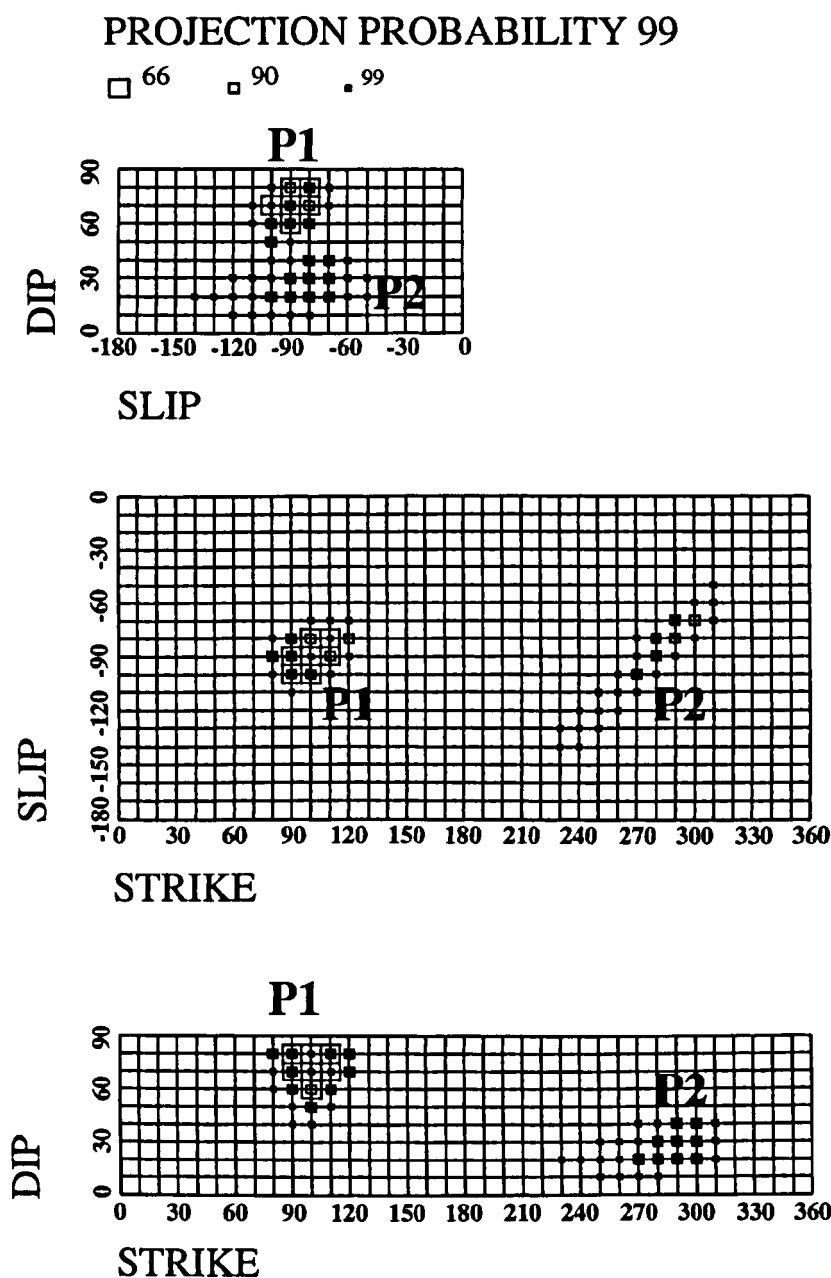


Figure 6. Result of the joint inversion of P polarities and S -wave polarization directions for event 491. Representation of the probability density function in the fault-plane parameter space (strike, dip, slip). The 66, 90 and 99 per cent confidence levels are represented with decreasing symbol sizes. The clusters correspond to the two conjugate fault planes P1 and P2.

by hand. This is related to the shape of the conditional probability density function used in the inversion. Because this function has a finite width around nodal planes, the algorithm naturally tends to favour solutions in which no station is very close to a nodal plane. In any case, variations in strike and dip between the two methods are small and within the uncertainties when both kinds of data (P and S) are well explained. In other cases, such as event 475, the mechanism, which was not well constrained in the manual solution, changes to a N-S- rather than a NW-SE-trending normal fault. Finally, events 92 and 97 (Fig. 5), which were badly constrained by the P polarities alone, are now better resolved.

We determined 148 focal mechanisms (Appendices C and

D). In the interpretation, we always choose the solution from the joint inversion, if it exists.

3 EARTHQUAKE AND FOCAL MECHANISM DISTRIBUTION

The seismicity appears to be located beneath the Gulf and a few kilometres inland on both sides of the Gulf (Fig. 8). More activity is observed on the northern side. In addition to some distributed background seismicity, three main clusters are identified: the cluster near the city of Aigion (C1), which was very active during the 2 months of the experiment and should be associated with the $M_L=4.5$ event that occurred there on

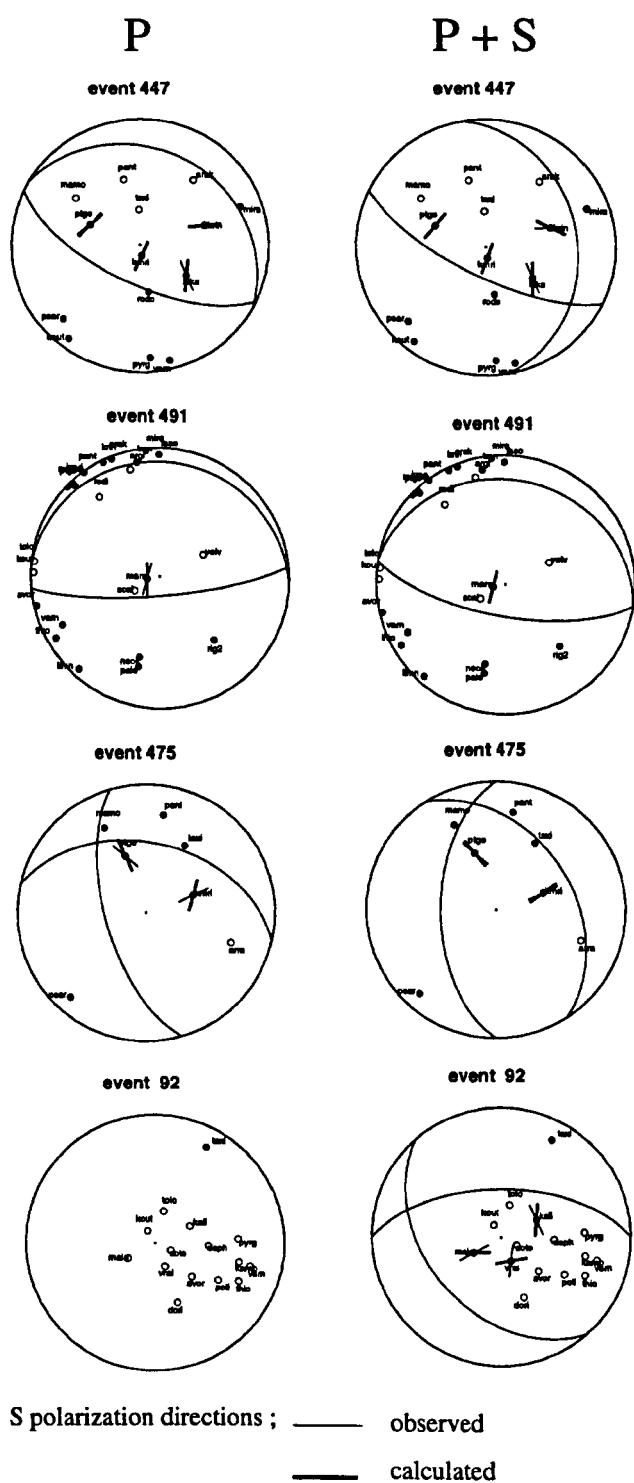


Figure 7. Comparison between focal mechanisms determined with *P* polarities only (left) and the focal mechanisms determined by joint inversion of *P* polarities and *S*-wave polarization directions (right). Thick lines are observed shear-wave polarizations and thin lines are calculated ones.

July 3; cluster Cl2, active during the month of August, with two $M_L=2.8-3.0$ events (47 and 677); and the cluster of August 24 (Cl3) with magnitude 2.0–2.8 events. The largest recorded events, both of magnitude $M_L = 3.0$, are the August 5

event (156 in Appendix B) and the August 23 event (677 in Appendix B).

The depth distribution of the events (Fig. 9) indicates that the maximum of activity is concentrated at 6–10 km depth. Moreover, 98 per cent of the events have focal depths less than 13–14 km. We deduce that the thickness of the seismogenic layer in the Patras–Aigion area does not exceed 15 km, considering the focal depth uncertainties of ± 2 km. This is in agreement with the aftershock study of the 1981 Corinth earthquakes by King *et al.* (1985), who found fewer than five events deeper than 15 km in the Corinth–Perachora region among a set of 133 well-located events. Another aspect of this vertical distribution is the absence of events shallower than 4 km (Fig. 9). We checked that this was not an artefact of our selection criteria for events. This small shallow activity is probably related to the rheological property of the shallow crustal layers, as suggested by King *et al.* (1994) for the San Andreas fault zone.

The distribution of the seismicity does not outline the geometry of the active faults at depth (Fig. 10). In the western part of the network (a1 and a2 in Fig. 10), the events align along an apparent geometry dipping $10^\circ-20^\circ$ north between 7 and 12 km depth, including cluster Cl2. This geometry does not appear clearly in the eastern sections (b and c in Fig. 10), except if we consider that it is outlined by the lower limit of the seismicity. Nevertheless, clusters Cl2 and Cl3 are in the down-dip extension of the Pspathopyrgos fault (P in Fig. 10) and the Aigion fault (Ai in Fig. 10), respectively.

The fault-plane solutions yield a consistent N–S direction of extension, in agreement with the large-scale tectonics of this area (Fig. 11). The majority of the mechanisms correspond to E–W-trending normal faulting consistent with the strike of the active faults. Few strike-slip and reverse fault-plane solutions are observed (e.g. 8, 86, 294, 308, and 358 in Appendix C).

Normal-faulting mechanisms with T axes oriented N–S and nodal planes dipping north or south between 35° and 70° are found in the eastern part of the network (Fig. 11a). The mean T axes for clusters Cl1 and Cl3 are subhorizontal in both cases (Table 2).

Normal-faulting mechanisms with a N–S-oriented T axis plunging approximately 30° are found in the western part of the network (Fig. 11b). They represent normal faulting with an E–W nodal plane dipping $75^\circ-85^\circ$ south and the other nodal plane dipping $10^\circ-25^\circ$ north. Most of these mechanisms are well constrained by *S*-wave polarization directions. These peculiar fault-plane solutions are found in cluster Cl2 but also south and northwest of it.

4 DISCUSSION

To assess the significance of microearthquakes within a volume of crustal rocks with major active faults is a difficult problem.

Table 2. Mean *P* and T axes for each cluster deduced from fault-plane solutions. Uncertainties are 10° for the strike and 5° for the plunge.

Cluster	Mean P Axis		Mean T Axis	
	Strike ($^\circ$)	Plunge ($^\circ$)	Strike ($^\circ$)	Plunge ($^\circ$)
Cl1	103.4	78.8	184.9	13.0
Cl2	354.9	66.0	187.5	29.8
Cl3	179.7	70.8	341.4	23.8

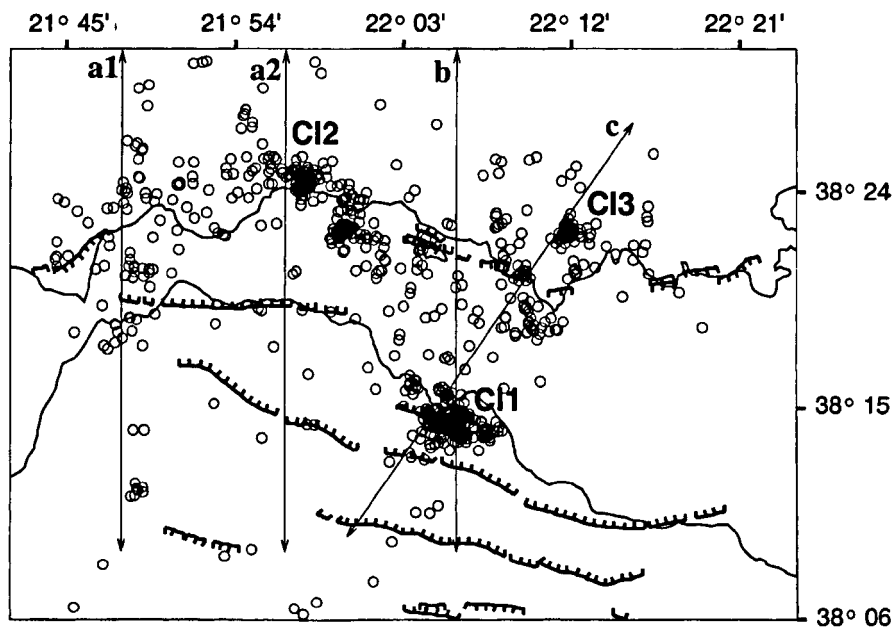


Figure 8. Locations of the 774 selected events. Segments a1, a2, b and c correspond to the sections shown in Fig. 11. CI1, CI2 and CI3 indicate the three main clusters.

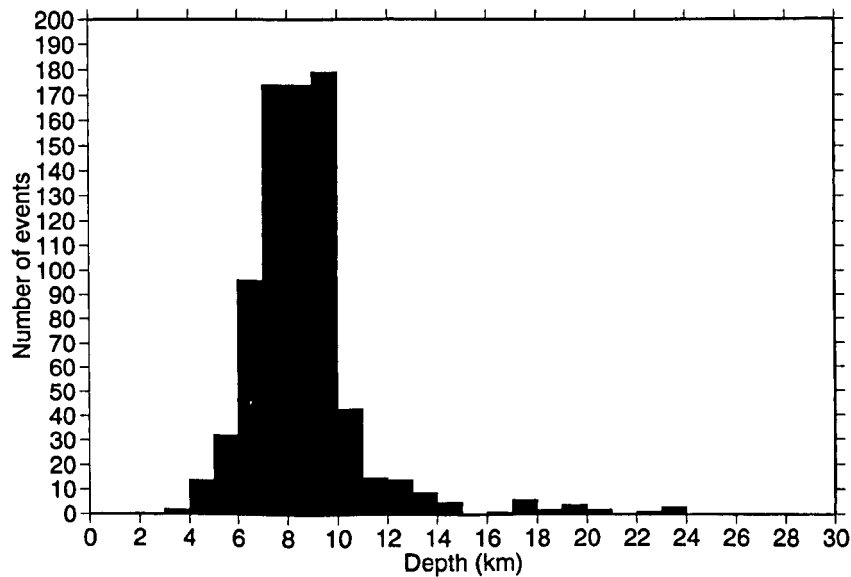


Figure 9. Vertical distribution of the 774 selected events. Uncertainties on focal depths are ± 2 km. 80 per cent of the activity is located between 6 and 10 km depth.

An event of magnitude smaller than 3 is not necessarily associated with slip on a major fault. Such an event corresponds to a source radius smaller than 300–400 m with only a few millimetres of slip. Thus, microearthquakes may be associated with major fault planes but also with minor faults. The presence of fractures or fracture zones in the upper crust is commonly assumed in active tectonic regions, where they are often observed near fault planes (Chester & Logan 1986; Scholz 1990). During the interseismic period, stress accumulation and elastic deformation of the crust around major fault planes can induce microseismic activity within rock volumes of the order of a few tens of cubic kilometres. Zones where microseismic activity is likely to be concentrated are the process zones near the extremities of major fault planes where stresses accumulate

(Cottrell 1964), or highly fractured regions located above ductile shear zones. Microseismicity may also represent internal deformation of blocks bounded by large faults (Fuchs *et al.* 1987; Jackson & White 1989; Echtler, Lüchen & Mayer 1994).

If the occurrence of microevents is controlled by a pre-existing network of fractures near major fault planes, where stresses may be locally modified, then some focal mechanisms may not be consistent with those expected from the large-scale tectonics. This is the case for the reverse and the few strike-slip events observed (e.g. events 8, 39, 79, 86, 294; Appendices C and D). All of these events are compatible with minimum horizontal (compressive) stress oriented N–S. However, modification of stresses may also be transient, for instance in response to slip episodes on major faults (e.g. Lyon-Caen *et al.*

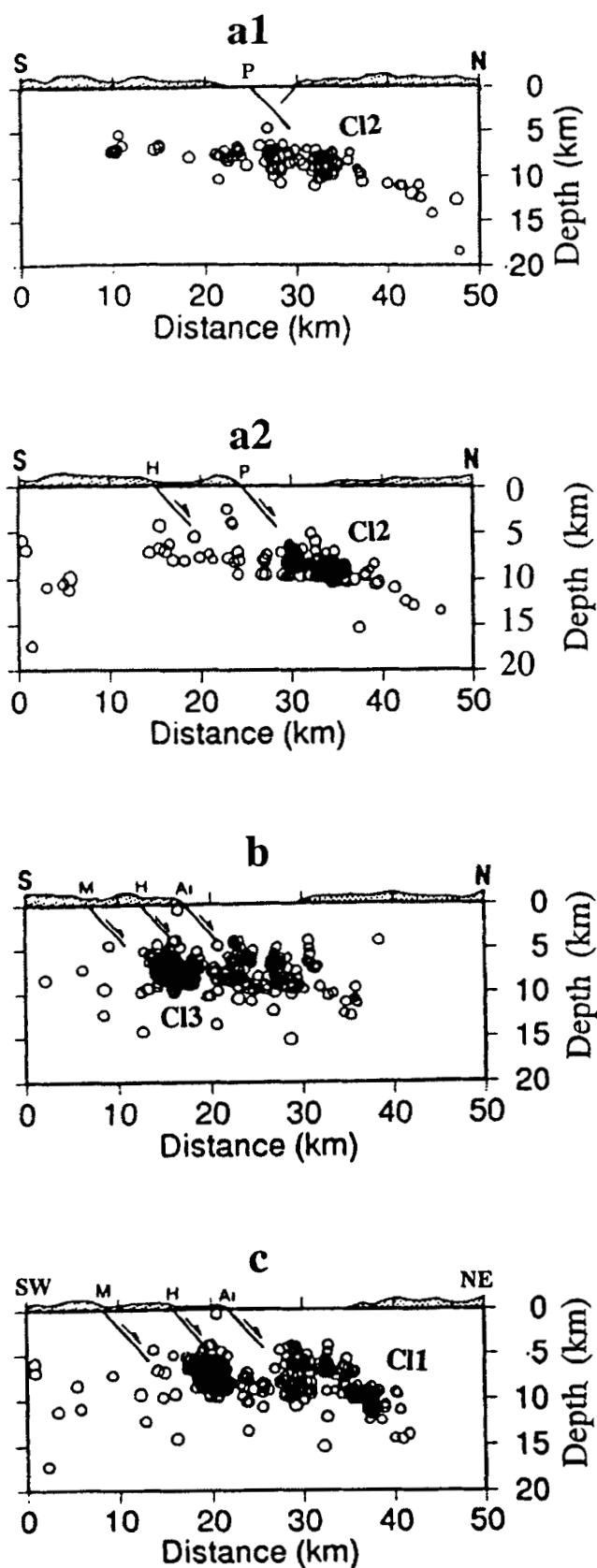


Figure 10. Vertical sections of seismicity (see Fig. 8 for location), with no vertical exaggeration. The half-width of projection is 6 km. Active faults observed at the surface have been extrapolated to a depth of 5 km with a dip of 50°. P: Psathopyrgos fault; H: Helike fault; Ai: Aigion fault; M: Mamousia fault.

© 1996 RAS, *GJI* 126, 663–688

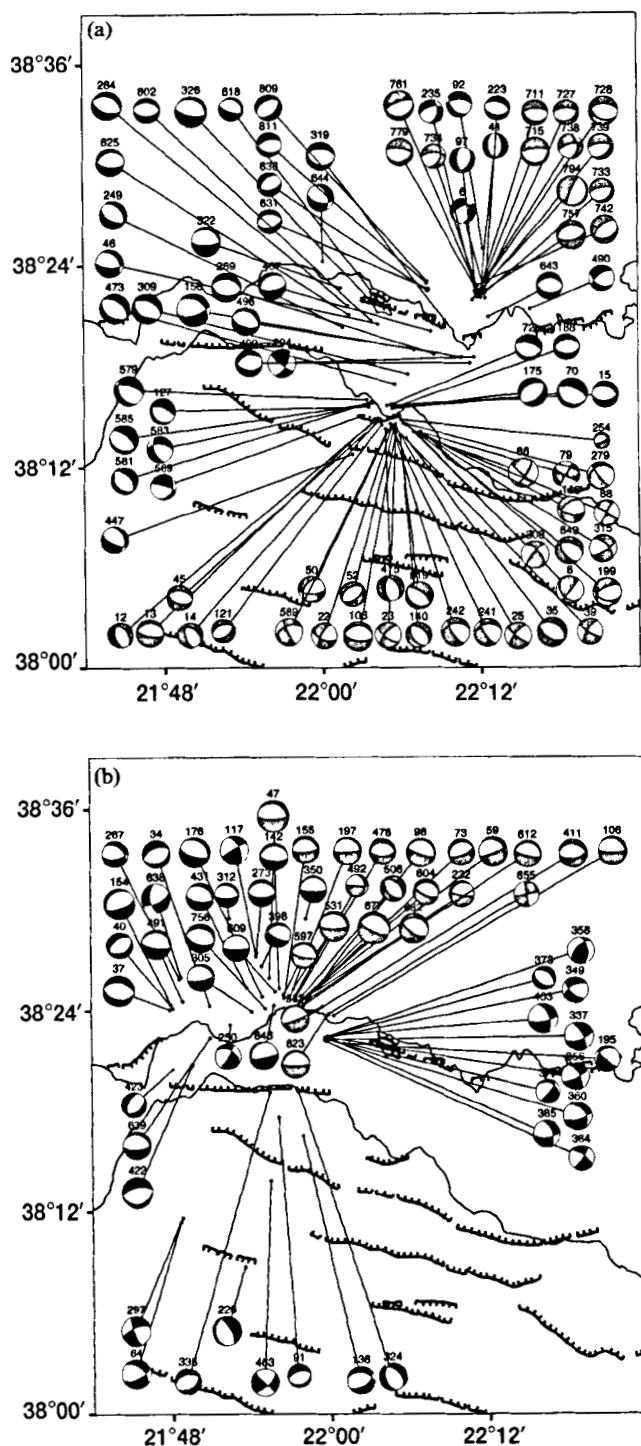


Figure 11. Maps with the 148 calculated fault-plane solutions. The size of the sphere (lower hemisphere) is proportional to magnitude. Grey fault-plane solutions are those within the three main clusters C11, C12 and C13. (a) Focal mechanisms in the western part of the network. (b) Focal mechanisms in the eastern part of the network. Numbers refer to events in Appendices B and C.

1988). This may be the case for cluster C11, which represents aftershocks of the $M_L = 4.5$ earthquake of July 3, 1991.

The two sections in Fig. 12 show the observed seismicity and some representative fault-plane solutions, with the geological information. In the western part of the network

(section a, Fig. 12) the distribution of the seismicity with depth increasing towards the north is consistent with the gently north-dipping nodal planes of the focal mechanisms (events 312, 491, 639, 677) and with the mean P and T axes of cluster C12. A likely interpretation is that a low-angle normal fault or a detachment zone occurs between 9 and 11 km depth, dipping north at about 15°. The seismicity could also be associated with slip on small antithetic faults corresponding to the nodal planes dipping steeply to the south. However, a network of such antithetic faults with the geometry given by the observed seismicity distribution may also accommodate overall strain compatible with the proposed detachment zone. If the dip of the Pspathopyrgos fault decreases from 50° as observed at the surface to less steep values at depth, then it would connect with the detachment in the region of cluster C12. The intersec-

tion of the two fault zones would be a likely place for increased fracturing and thus for generating microseismic activity. The possible upward and downward extensions of the proposed detachment are difficult to define due to the lack of observed microevents. Nevertheless, it seems likely that the detachment extends down-dip to the north, dying out in the ductile lower part of the crust, where it would become aseismic.

The relationship between the spatial event distribution, the focal mechanisms and the surface geology in the eastern part of the network (section b, Fig. 12) appears to be more complex than to the west. The occurrence of a detachment on section b can only be justified by the geometry of the observed seismicity and would correspond to the lower limit of the seismogenic layer. This lower limit appears to dip gently northwards, although no determined fault-plane solution is

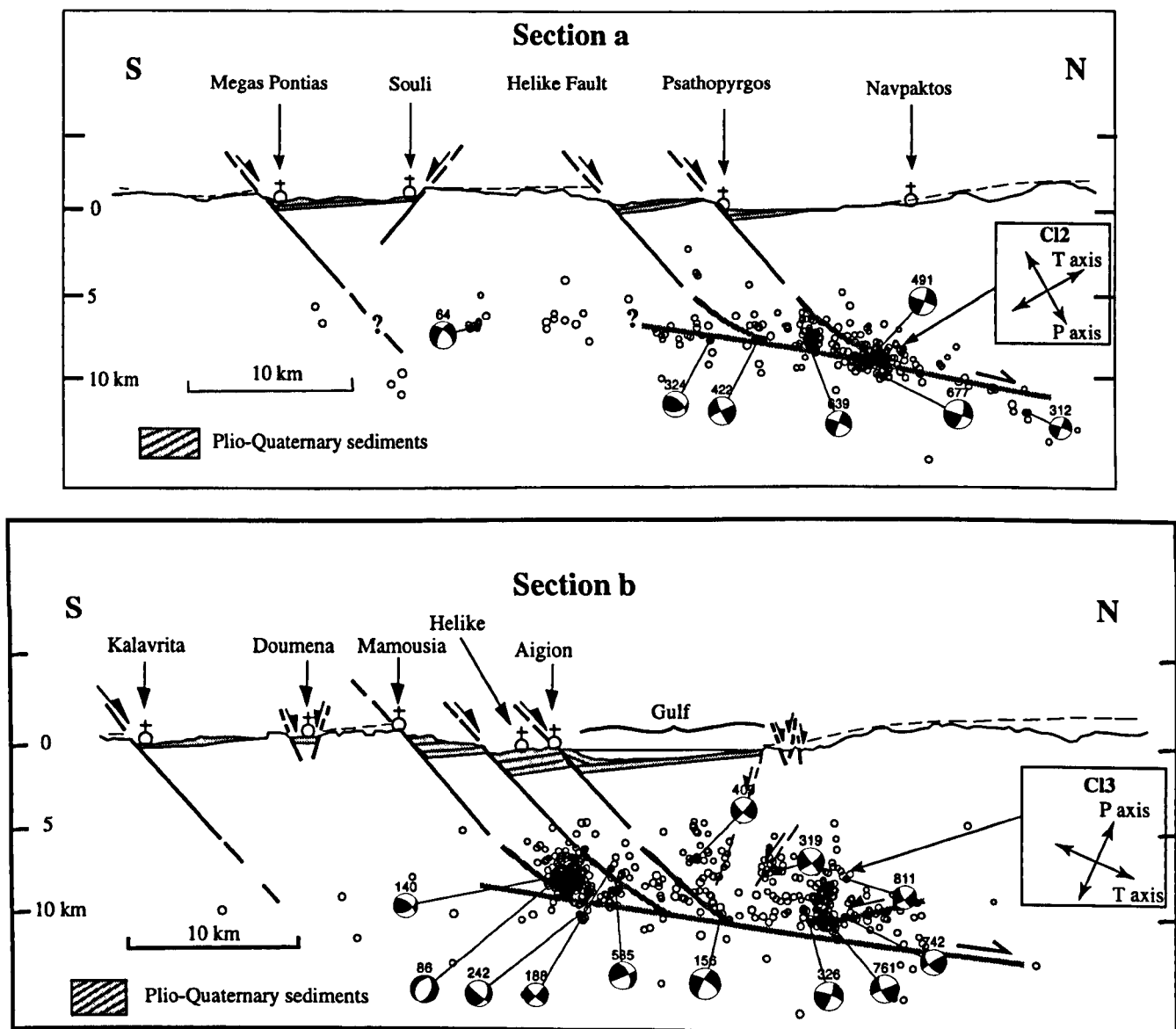


Figure 12. Interpretative sections (locations in Fig. 1). Only a few representative fault-plane solutions have been projected on the vertical plane. Insets represent mean P and T axes for clusters C12 and C13. The active major normal faults documented at the surface connect with the northward-dipping plane (shaded), consistent with the spatial distribution of seismicity and the gently north-dipping nodal planes of the western focal mechanisms.

clearly associated with it. As in section a, the major active normal faults documented at the surface (the Mamousia, Helike and Aigion faults) may connect with the hypothetical detachment at depth. Small antithetic faults probably related to flexure of the hanging wall have been recognized in the field (Fig. 1). Because the total amount of N–S extension is greater in section b than in section a, the hanging wall (northern part of the Gulf) may also be more fractured (Rigo 1994; Armijo *et al.* 1996). Part of the observed shallow (0–15 km) seismicity in the hanging wall of the Helike–Aigion faults could be associated with such flexure-related antithetic faulting, in agreement with the 60°–65° south-dipping nodal planes of some mechanisms (dashed lines in section b, Fig. 12). Farther north in the hanging wall, cluster C13 is located at an average depth of about 10 km. The mean P axis of cluster C13 is steeper than those of clusters C11 and C12 (Figs 12, 10 and 9). Also, unlike the other two clusters, the mean T axis of cluster C13 plunges to the north (Table 2; Fig. 12, section b). Cluster C13 may thus be associated with gently south-dipping faulting above the detachment. Finally, the Aigion cluster (C11, Fig. 9) is located roughly in the region where the downward extension of the Mamousia fault may connect with the proposed detachment (Fig. 12). However, the seismicity associated with this cluster appears to diffuse over almost all of the lower part of the block bounded by the Mamousia and the Helike faults.

5 CONCLUSIONS

Because of the high resolution of the data presented here, we can provide a precise, instantaneous image of the deforming study area. This high resolution enables us to describe precisely the mechanisms of small earthquakes and to draw some inferences about the large-scale geometry and mechanics of extension in the Gulf of Corinth. The data set suggests the existence, beneath the Gulf of Corinth, of a detachment zone that contributes significantly to the overall deformation process. All the large active normal faults mapped appear to connect downwards with the proposed detachment at around 10 km depth. The geometry of the detachment would be that of a normal fault dipping 15° ($\pm 10^\circ$) northwards, as revealed by the studied seismicity. This is similar to the geometry and mechanism suggested by King *et al.* (1985) for the eastern part of the Corinth Gulf in the Corinth–Perachora region.

Except for cluster C12, the observed cluster activity does not exhibit a simple geometry associated with any particular fault. Fault-plane solutions can be quite varied within one cluster, as in the Aigion cluster (C11). The cluster microearthquake activity is likely to be mostly located on small structures, antithetic faults or fracture zones located near the extremities of the large faults.

To evaluate roughly the total seismic moment released by the 774 well-located events in the network, we use the formula relating the magnitude M_L to the seismic moment M_0 established in the Volos region in Central Greece ($\log M_0 = 1.5 M_L + 9.0$; Makropoulos, personal communication). The total seismic moment released during the 30 day period is of the order of $0.3\text{--}3.0 \times 10^{16}$ N m, which corresponds to an event with an M_s of 4.3–4.9. This represents quite a high rate of diffuse deformation which is unlikely to be a permanent feature. Previous seismicity experiments in the area indicate activity beneath the Gulf between Patras and Aigion (Melis *et al.* 1989) and on the northern side of the Gulf (Hatzfeld *et al.* 1990).

Because the resolution of these studies is much poorer than the one presented here, it is difficult to make valid comparisons. However, it appears that the Aigion cluster C11 was not active at the time of these experiments (1983–84 and 6 weeks during 1986). Thus, the cluster activity probably changes with time, contributing to the redistribution of stress at some depth. In the Gulf of Corinth, the focal depths of large earthquakes occur near the base of the seismogenic layer at 10–15 km depth, as in other regions under extension throughout the world (Sibson 1982; Meissner & Strehlau 1982; Jackson & White 1989). Therefore, we suggest that the microseismic activity observed corresponds to episodes of stress transfer from the lower semi-brittle zone to the upper brittle zone, through slip on a detachment zone. This process may cause stresses to concentrate in regions near the base of the large steep faults seen at the surface. Such regions are likely places for large normal-faulting earthquakes to initiate their rupture. However, our observations do not tell us if this activity represents to some extent the initiation of a preseismic stage. That the microseismic activity is most probably transient suggests that the stress accumulation near the large faults may operate in cycles with small spatial extent and short duration which may only correspond to discrete episodes during the main seismic cycle.

ACKNOWLEDGMENTS

The authors thank all the participants of the summer 1991 field experiment. We also thank Geneviève Patau for reading some of the seismograms, Bertrand Meyer for helpful discussions and Kurt Feigl for a helpful reading of manuscript. The Direction aux Risques Majeurs (DRM) and Institut National des Sciences de l'Univers (INSU) provided financial support. We also benefited from EEC EPOCH Programme support. This is IPGP contribution no. 1427.

REFERENCES

- Abers, G., 1991. Possible seismogenic shallow-dipping normal faults in the Woodlark–D'entrecasteaux extensional province, Papua New Guinea, *Geology*, **19**, 1205–1208.
- Ambraseys, N.N. & Jackson, J.A., 1990. Seismicity and associated strain of Central Greece between 1890 and 1988, *Geophys. J. Int.*, **101**, 663–708.
- Amorese, D., 1993. Sismotectonique et déformation actuelle de la terminaison nord occidentale de l'Arc Égéen (Iles Ioniennes, Acharnanie, Épire, Grèce), *Thèse de Doctorat de 3ème cycle*, Université Joseph Fourier, Grenoble, France.
- Arabas, W.J. & Julander, D.R., 1986. Geometry of seismically active faults and crustal deformation within the Basin and Range—Colorado Plateau transition in Utah, *Spec. Pap. Geol. Soc. Am.*, **208**, 43–74.
- Armijo, R., Meyer, B., King, G.C.P., Rigo, A. & Papanastassiou, D., 1996. Quaternary evolution of the Corinth Rift and its implications for the Late Cenozoic evolution of the Aegean, *Geophys. J. Int.*, **126**, 11–53.
- Bernard, P. & Zollo, A., 1989a. The Irpinia (Italy) 1980 earthquake: detailed analysis of a complex normal fault, *J. geophys. Res.*, **94**, 1631–1647.
- Bernard, P. & Zollo, A., 1989b. Inversion of near source S polarization for parameters of double couple point sources, *Bull. seism. Soc. Am.*, **79**, 1779–1809.
- Booth, D.C. & Crampin, S., 1985. Shear wave polarizations on a

- curved wavefront at an isotropic free surface, *Geophys. J. R. astr. Soc.*, **88**, 31–45.
- Bouin, M.P., T  llez, J. & Bernard, P., 1996. Seismic anisotropy around the Gulf of Corinth, Greece, deduced from three-component seismograms of local earthquakes and its relationship with crustal strain, *J. geophys. Res.*, **101**, 5797–5811.
- Brillinger, D.R., Udias, A. & Bolt, B.A., 1980. A probability model for regional focal mechanism solutions, *Bull. seism. Soc. Am.*, **70**, 149–170.
- Briole, P., Deschamps, A., Lyon-Caen, H., Papazissi, K. & Martinod, J., 1993. The Itea ($M_s=5.9$) earthquake of November 18, 1992. Characteristics of the main shock inferred from body waves and ground displacement analysis, *Proc. 2nd Cong. of the Hellenic Geophys. Un. in Florina*, **1**, 297–308.
- Chester, F.M. & Logan, J.M., 1986. Implications for mechanical properties of brittle faults from observations of the Punchbowl fault zone, California, *Pageoph.*, **124**, 79–106.
- Cottrell, A.H., 1964. *The Mechanical Properties of Materials*, John Wiley and Sons, New York, NY.
- Covington, H.R., 1983. Structural evolution of the Raft River Basin, Idaho, *Geol. Soc. Am. Memoir*, **157**, 229–237.
- De Chabali  r, J.B., Lyon-Caen, H., Zollo, A., Deschamps, A., Bernard, P. & Hatzfeld, D., 1992. A detailed analysis of microearthquakes in western Crete from digital three-component seismograms, *Geophys. J. Int.*, **110**, 347–360.
- Echtler, H.P., L  schen, E. & Mayer, G., 1994. Lower crustal thinning in the Rhinegraben: Implications for recent rifting, *Tectonics*, **13**, 342–353.
- Evans, R., 1984. Effects of the free surface on shear wavetrains, *Geophys. J. R. astr. Soc.*, **76**, 165–172.
- Eyidogan, H. & Jackson, J., 1985. A seismological study of normal faulting in the Demirci, Alaschir and Gediz earthquakes of 1969–70 in western Turkey: implications for the nature and geometry of deformation in the continental crust, *Geophys. J. R. astr. Soc.*, **81**, 569–607.
- Fuchs, K., Bonjer, K.-P., Gajewski, D., L  schen, E., Prodehl, C., Sandmeier, K.-J., Wenzel, F. & Wilhelm, H., 1987. Crustal evolution of the Rhinegraben area. 1. Exploring the lower crust in the Rhinegraben rift by unified geophysical experiments, *Tectonophysics*, **141**, 261–275.
- Galanopoulos, A.G., 1960. Greece, a catalogue of shocks with $I_0 \geq VI$ or $M \geq 5$ for the years 1801–1958, Academy of Science, Athens.
- Galanopoulos, A.G., 1961. Greece, a catalogue of shocks with $I_0 \geq VII$ for the years prior to 1800, Academy of Science, Athens.
- Hatzfeld, D., Pedotti, G., Hatzidimitriou, P. & Makropoulos, K., 1990. The strain pattern in the Western Hellenic Arc deduced from a microearthquake survey, *Geophys. J. Int.*, **101**, 181–202.
- Jackson, J.A. & White, N.J., 1989. Normal faulting in the upper continental crust: observations from regions of active extension, *J. Struct. Geol.*, **11**, 15–36.
- Jackson, J.A., Gagnepain, J., Houseman, G., King, G.C.P., Papadimitriou, P., Souferis, C. & Virieux, J., 1982. Seismicity, normal faulting and the geomorphological development of the Gulf of Corinth (Greece): the Corinth earthquakes of February and March 1981, *Earth planet. Sci. Lett.*, **57**, 377–397.
- Jones, H.C., 1987. A geophysical and geological investigation of extensional structures, Great Basin, Western United States, *PhD thesis*, Massachusetts Institute of Technology, Cambridge, MA.
- King, G.C.P., Ouyang, Z.X., Papadimitriou, P., Deschamps, A., Gagnepain, J., Houseman, G., Jackson, J.A., Souferis, C. & Virieux, J., 1985. The evolution of the Gulf of Corinth (Greece): an aftershock study of the 1981 earthquakes, *Geophys. J. R. astr. Soc.*, **80**, 677–683.
- King, G.C.P., Oppenheimer, D. & Amelung, F., 1994. Block versus continuum deformation in the western United States, *Earth planet. Sci. Lett.*, **128**, 55–64.
- Lee, W.H.K. & Lahr, J.C., 1975. HYPO71 (revised): a computer program for determining hypocenter, magnitude and first motion pattern of local earthquakes, *US Geol. Survey Open-File report*, **75–311**, 116 pp.
- Le Meur, H., 1994. Tomographie tridimensionnelle    partir des temps des premi  res arriv  es des ondes P et S, application    la r  gion de Patras, *Th  se de Doctorat de 3  me cycle*, Universit   Paris 7, France.
- Logatchev, N.A. & Zorin, Y.A., 1987. Evidence and causes of the two-stages development of the Baikal rift, *Tectonophysics*, **143**, 225–234.
- Lyon-Caen, H., Armijo, R., Drakopoulos, J., Baskoutas, J., Delibassis, N., Gaulon, R., Kouskouna, V., Latoussakis, J., Makropoulos, K., Papadimitriou, P., Papanastassiou, D., Pedotti, G., 1988. The 1986 Kalamata (South Peloponnesus) earthquake: detailed study of a normal fault and tectonic implications, *J. geophys. Res.*, **93**, 14 967–15 000.
- Maier, L. & Eisbacher, G.H., 1991. Crustal kinematics and deep structure of the northern Rhine Graben, Germany, *Tectonics*, **10** (3), 621–630.
- Meissner, R. & Strehlau, J., 1982. Limits of stress in continental crust and their relation to the depth–frequency relation of shallow earthquakes, *Tectonics*, **1**, 73–89.
- Melis, N.S., Brooks, M. & Pearce, R.G., 1989. A microearthquake study in the Gulf of Patras region, western Greece, and its seismotectonic interpretation, *Geophys. J. R. astr. Soc.*, **98**, 515–524.
- Papazachos, B. & Papazachos, K., 1989. *The Earthquakes in Greece*, Thessaloniki, Ziti Publications, Greece (in Greek).
- Rigo, A., 1994.   tude sismotectonique et g  od  sique du Golfe de Corinthe (Gr  ce), *Th  se de Doctorat de 3  me cycle*, Universit   Paris 7, France.
- Schmidt, J.F.J., 1881. *Studien   ber Vulkans und Erdbeben*, C. Scholtze, Leipzig (in German).
- Scholz, C.H., 1990. *The Mechanics of Earthquakes and Faulting*, Cambridge University Press, Cambridge.
- Sibson, R.H., 1982. Fault zone models heat flow and depth distribution of earthquakes in the continental crust of the United States, *Bull. seism. Soc. Am.*, **72**, 151–163.
- Taymaz, T., Jackson, J. & McKenzie, D., 1991. Active tectonics of the north and central Aegean Sea, *Geophys. J. Int.*, **106**, 433–490.
- Wenzel, F., Brun, J.P. & ECORS-DEKORP working group, 1991. A deep reflection seismic line across the northern Rhine Graben, *Earth planet. Sci. Lett.*, **104**, 140–150.
- Zollo, A. & Bernard, P., 1991. Fault mechanisms from near source data: joint inversion of S polarizations and P polarities, *Geophys. J. Int.*, **104**, 441–451.

**APPENDIX A: COORDINATES OF STATIONS
DURING THE 1991 EXPERIMENT**

Sites	latitude (°N)	longitude (°E)	Sites	latitude (°N)	longitude (°E)
daph	38°25.27'	22°06.21'	marm	38°24.66'	21°49.71'
pyrg	38°24.64'	22°00.99'	pale	38°30.02'	21°50.70'
kout	38°24.04'	22°13.07'	eupa	38°26.13'	21°55.36'
varn	38°27.88'	21°57.58'	iaso	38°13.68'	21°48.94'
trik	38°26.76'	21°54.20'	mira	38°09.31'	21°51.10'
triz	38°21.79'	22°04.30'	arra	38°15.96'	21°58.96'
scal	38°25.24'	21°50.48'	piti	38°15.97'	21°53.90'
neok	38°29.30'	21°50.48'	anoz	38°18.76'	21°56.89'
velv	38°23.38'	21°45.98'	arak	38°08.10'	21°57.44'
avor	38°28.04'	22°08.23'	liko	38°13.44'	21°56.60'
rig1	38°28.83'	21°46.05'	mamo	38°09.41'	22°07.92'
rig2	38°28.21'	21°44.84'	taxi	38°10.71'	22°02.42'
thio	38°29.52'	22°00.08'	kefa	38°11.24'	21°49.59'
tolo	38°22.41'	22°11.37'	tran	38°15.74'	21°51.12'
psar	38°19.51'	22°11.07'	kmri	38°13.42'	22°02.23'
limn	38°32.45'	21°58.01'	pant	38°08.48'	22°03.67'
dori	38°30.81'	22°09.61'	ptge	38°11.47'	22°06.22'
pano	38°22.97'	22°16.30'	krin	38°11.54'	21°57.11'
vrai	38°26.99'	22°11.24'	derv	38°10.31'	22°08.24'
poti	38°28.77'	22°04.43'	agge	38°12.97'	22°05.20'
palp	38°28.77'	21°52.70'	rodo	38°15.65'	22°01.71'
pits	38°25.93'	21°44.89'	akas	38°16.62'	21°50.76'
krok	38°31.87'	22°04.27'	dafn	38°12.49'	22°00.30'
mala	38°26.38'	22°15.16'	fter	38°09.36'	22°04.61'
elai	38°23.36'	22°07.08'	argi	38°16.96'	21°51.56'
kmro	38°24.70'	21°56.93'	univ	38°17.32'	21°47.33'
klim	38°25.21'	21°58.09'	rodi	38°19.42'	21°53.87'
palr	38°27.78'	21°52.23'	laka	38°14.67'	21°58.76'
sote	38°25.66'	22°10.71'	koul	38°14.33'	22°05.18'
kamb	38°27.20'	22°00.54'	koun	38°11.19'	22°00.72'
kali	38°23.61'	22°08.45'	mal1	38°22.37'	21°52.68'
serg	38°24.82'	22°03.43'	mal2	38°22.81'	21°54.22'

APPENDIX B: EVENT LOCATIONS FROM HYPOT1.

Time origin time (h mn s); Lat.: latitude N (°mn); Long.: longitude E (°mn); depth in km; ERZ: precision on the focal depth in km; Mag.: Magnitude Mi; RMS: Root Mean Square of the computed location in sec; ERX and ERY precision on the epicentral position in km.

Table with 20 columns: Date, Time, Lat, Long, Depth ERZ, Mag., RMS, ERX, ERY, n.evt. It contains two identical data blocks, one for Date 910802 and one for Date 910801, listing seismic event details.

APPENDIX B (Continued.)

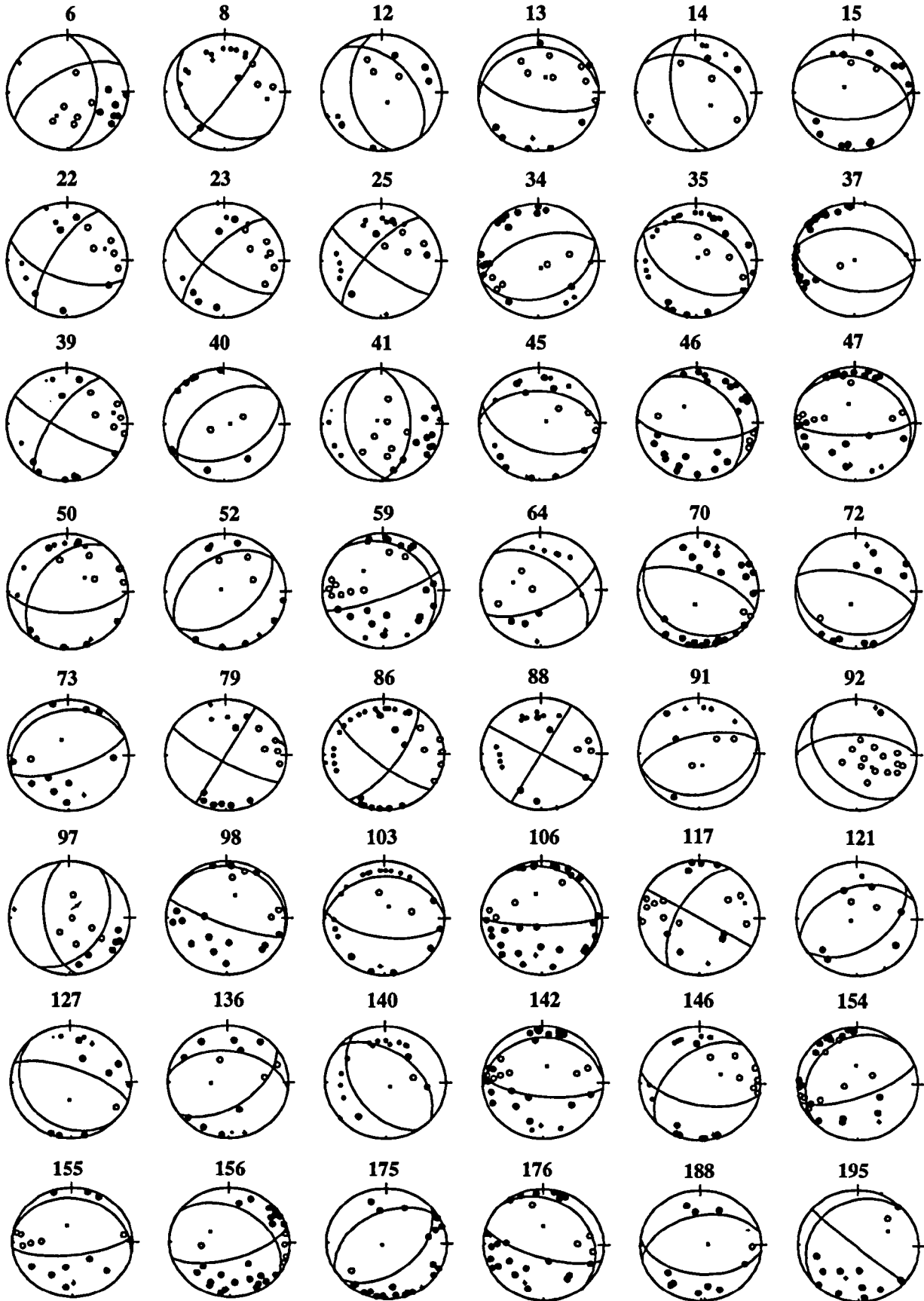
Date	Time	Lat	Long	Depth ERZ	Mag.	RMS	ERY	n.evt	
910812	0 45 1.66	38 22.38	21 59.78	7.56	0.4	2.44	0.10	0.2	355
910812	0 47 45.08	38 22.38	21 59.77	8.55	0.3	2.59	0.11	0.1	356
910812	0 55 46.56	38 15.65	22 5.66	8.00	0.4	2.38	0.08	0.3	357
910812	1 22 9.25	38 22.39	21 59.81	7.85	0.3	2.57	0.10	0.1	358
910812	1 39 44.15	38 22.31	21 59.75	7.78	0.4	2.42	0.11	0.2	359
910812	1 48 47.36	38 22.26	21 59.87	8.17	0.2	2.64	0.11	0.1	360
910812	3 85 9.34	38 23.25	22 0.05	9.52	1.4	2.65	0.23	0.6	361
910812	3 20 31.93	38 22.46	21 59.73	6.85	0.4	2.40	0.10	0.2	362
910812	3 53 45.19	38 13.15	22 3.40	9.47	0.6	2.54	0.16	0.3	363
910812	4 26 3.51	38 22.16	21 59.74	7.47	0.5	2.40	0.15	0.3	364
910812	4 57 8.24	38 21.17	22 2.04	8.55	0.3	2.51	0.07	0.1	365
910812	5 73 8.19	38 22.67	21 46.15	8.65	0.3	2.83	0.11	0.2	366
910812	5 13 7.68	38 24.80	21 57.21	8.89	0.4	2.37	0.10	0.2	367
910812	5 29 27.39	38 21.78	22 15.95	12.08	0.5	2.52	0.17	0.6	368
910812	6 02 0.35	38 22.63	21 59.96	8.21	0.8	2.38	0.14	0.4	369
910812	6 13 51.81	38 22.44	21 59.72	8.05	0.3	2.37	0.13	0.2	370
910812	6 47 48.54	38 17.23	22 4.40	9.69	0.9	2.46	0.17	0.4	371
910812	7 25 4.12	38 22.16	22 1.58	8.78	0.3	2.42	0.08	0.2	372
910812	7 55 7.39	38 22.38	21 59.64	8.57	0.5	2.20	0.04	0.2	373
910812	7 59 17.59	38 14.24	22 5.96	7.42	0.5	2.46	0.13	0.4	374
910812	9 34 3.44	38 22.47	22 0.18	7.92	0.9	0.00	0.12	0.4	375
910812	9 19 42.10	38 22.49	21 59.80	7.50	0.5	0.00	0.08	0.2	376
910812	9 50 45.48	38 21.98	22 1.46	8.38	0.3	2.32	0.09	0.2	377
910812	9 52 56.67	38 25.79	21 49.06	9.13	0.5	2.14	0.09	0.3	378
910812	11 45 2.58	38 22.66	21 59.75	7.24	0.4	0.00	0.08	0.3	379
910812	13 35 55.73	38 14.45	22 7.25	7.58	0.5	2.52	0.11	0.3	380
910812	13 36 50.81	38 35.75	22 8.68	12.44	0.9	2.20	0.05	0.3	381
910812	14 32 44.75	38 14.73	22 6.22	7.40	0.4	2.69	0.12	0.2	382
910812	15 33 15.19	38 23.34	22 0.60	9.57	0.3	2.69	0.09	0.1	383
910812	15 54 57.03	38 23.91	21 48.24	8.12	0.3	2.65	0.10	0.2	384
910812	16 44 36.87	38 22.40	21 59.79	8.47	0.3	2.57	0.08	0.1	385
910812	16 58 28.49	38 18.26	22 10.03	6.13	0.5	2.35	0.10	0.6	386
910812	17 33 5.74	38 16.70	22 1.15	7.77	0.5	2.39	0.15	0.4	387
910812	17 81 3.00	38 20.01	22 5.85	9.08	0.6	2.57	0.17	0.3	388
910812	17 16 37.93	38 22.24	22 13.03	11.64	0.7	2.22	0.14	0.9	389
910812	18 22 45.59	38 22.34	21 59.31	6.82	0.4	0.00	0.04	0.2	390
910812	18 27 0.18	38 22.38	21 59.48	6.79	0.7	2.20	0.09	0.3	391
910812	18 28 29.71	38 20.54	22 9.42	6.40	0.3	2.25	0.08	0.2	392
910812	19 05 1.69	38 20.92	22 9.58	6.78	0.2	2.60	0.13	0.3	393
910812	20 17 50.04	38 20.55	22 9.42	6.44	0.4	2.31	0.12	0.3	394
910812	20 39 45.11	38 20.65	22 9.30	6.22	0.4	2.28	0.11	0.3	395
910812	20 55 13.37	38 7.63	22 1.42	10.90	0.7	2.50	0.12	0.4	396
910812	21 39 25.45	38 26.67	21 54.91	9.56	0.3	2.37	0.08	0.2	397
910812	22 03 5.12	38 22.65	21 59.82	7.45	0.5	2.34	0.12	0.3	398
910812	22 10 25.05	38 20.28	22 9.57	6.89	0.3	2.29	0.09	0.3	399
910812	23 28 19.44	38 33.03	21 52.88	17.28	0.4	2.11	0.05	0.5	400
910812	23 28 19.44	38 23.67	21 57.58	5.91	0.7	2.57	0.11	0.2	401
910812	23 35 32.51	38 22.46	21 59.87	8.50	0.4	2.84	0.14	0.2	402
910812	23 42 2.80	38 22.29	22 0.06	8.00	0.7	2.44	0.14	0.3	403
910813	0 52 5.03	38 22.52	21 59.73	7.46	0.6	0.00	0.08	0.2	404
910813	0 29 2.50	38 20.55	22 9.56	6.88	0.2	2.38	0.09	0.3	405

APPENDIX B (Continued.)

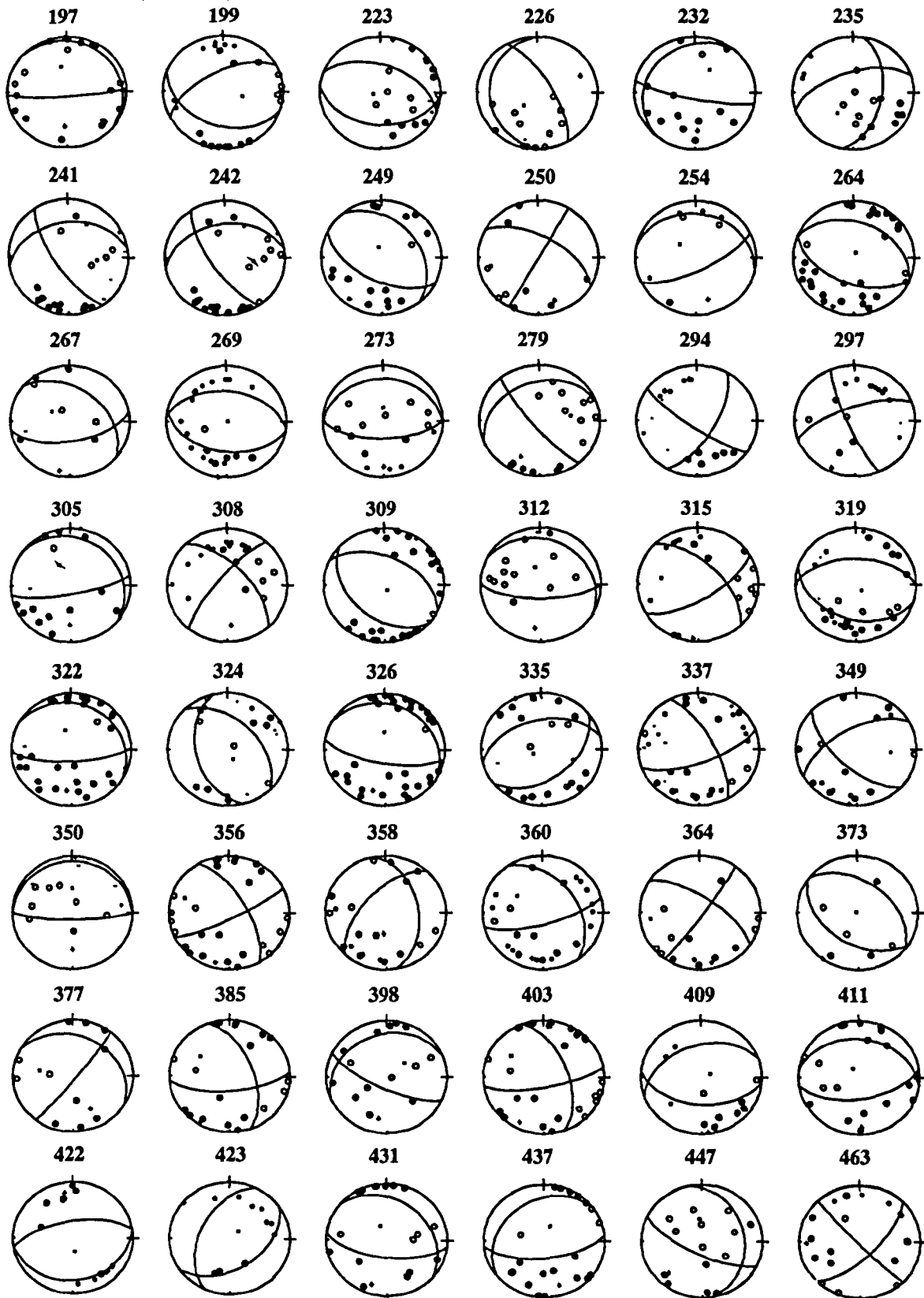
Date	Time	Lat	Long	Depth ERZ	Mag.	RMS	ERX	ERY	n.evt	Date	Time	Lat	Long	Depth ERZ	Mag.	RMS	ERX	ERY	n.evt	
910825	4 556.74	38 20.37	21 48.56	8.32	0.5	2.82	0.11	0.3	0.1	910826	1 4317.41	38 24.27	21 57.77	7.13	0.4	1.94	0.09	0.2	0.2	818
910825	4 4616.40	38 10.94	22 4.89	4.50	0.5	2.46	0.13	0.2	0.2	910826	1 5632.37	38 13.94	22 5.91	7.89	0.4	2.59	0.16	0.3	0.2	819
910825	5 1319.02	38 22.06	22 11.71	9.60	0.4	2.41	0.08	0.3	0.3	910826	2 743.46	38 25.24	22 8.27	10.61	0.9	2.25	0.11	0.6	0.8	820
910825	5 1621.73	38 22.35	22 11.84	9.67	0.4	2.84	0.13	0.3	0.2	910826	2 33 2.76	38 22.69	22 8.31	6.57	0.6	2.25	0.09	0.2	0.3	821
910825	5 5951.93	38 25.35	22 7.96	10.96	0.2	2.36	0.05	0.1	0.3	910826	3 813.45	38 16.14	22 3.73	11.97	2.3	0.00	0.14	0.7	0.5	822
910825	6 9 5.20	38 13.69	22 7.70	9.20	0.9	2.44	0.16	0.5	0.5	910826	4 149.52	38 14.11	22 5.46	8.00	0.4	2.53	0.16	0.3	0.2	823
910825	6 3651.11	38 14.54	22 5.29	7.53	0.8	2.21	0.14	0.5	0.5	910826	4 1423.63	38 22.65	22 0.01	7.88	0.4	1.86	0.05	0.2	0.2	824
910825	6 4940.60	38 14.42	22 5.74	7.86	0.8	2.14	0.15	0.4	0.4	910826	4 1434.31	38 23.09	22 0.35	8.62	0.5	2.06	0.11	0.3	0.3	825
910825	7 2254.70	38 13.86	22 4.70	6.94	0.4	2.36	0.07	0.3	0.5	910826	6 2140.32	38 18.49	21 52.98	7.80	0.5	2.75	0.11	0.3	0.2	826
910825	7 4056.40	38 22.40	22 11.81	9.37	0.3	2.39	0.10	0.3	0.2	910826	8 1826.91	38 28.99	21 58.59	12.28	0.4	2.49	0.11	0.3	0.3	827
910825	9 3349.37	38 21.81	22 11.68	9.86	0.4	1.96	0.04	0.5	0.2	910826	8 3755.60	38 18.35	22 18.99	10.46	0.9	2.76	0.11	0.7	0.4	828
910825	9 4724.14	38 20.76	22 3.78	9.15	0.5	2.46	0.08	0.2	0.1	910826	9 13 2.50	38 13.87	22 6.05	7.37	0.5	2.55	0.16	0.3	0.3	829
910825	11 2736.16	38 14.48	22 5.20	7.33	0.6	2.44	0.13	0.3	0.4	910826	9 2911.82	38 20.50	22 5.95	8.00	0.3	2.14	0.08	0.2	0.2	830
910825	11 30 2.97	38 13.86	22 5.80	7.54	0.5	2.49	0.15	0.3	0.3	910826	10 841.93	38 23.56	21 49.69	9.32	2.0	2.67	0.12	1.0	0.5	831
910825	13 2029.12	38 13.62	22 5.80	7.82	0.5	2.50	0.19	0.4	0.4	910826	11 3417.67	38 21.68	22 1.03	9.51	0.3	2.62	0.07	0.2	0.1	833
910825	14 3555.73	38 13.76	22 3.96	6.69	0.4	2.31	0.10	0.4	0.5	910826	16 1259.39	38 19.02	21 59.21	7.15	0.2	2.59	0.11	0.3	0.1	834
910825	18 4 5.00	38 23.87	21 49.81	7.67	0.3	2.16	0.07	0.2	0.4	910826	18 5614.29	38 24.96	21 55.12	10.18	0.4	2.50	0.04	0.4	0.3	835
910825	18 2518.62	38 24.63	21 54.74	9.32	0.3	2.64	0.10	0.1	0.2	910826	19 1450.34	38 14.17	22 5.52	5.98	0.9	2.35	0.14	0.6	0.4	836
910825	18 4956.44	38 23.08	22 8.01	6.91	0.3	2.48	0.13	0.2	0.4	910826	20 3311.15	38 24.61	21 58.94	8.55	0.4	2.44	0.08	0.3	0.3	837
910825	18 5155.10	38 22.97	22 7.98	7.19	0.2	2.34	0.11	0.2	0.3	910826	22 3928.52	38 22.08	22 0.40	6.79	0.7	2.46	0.12	0.4	0.3	838
910825	20 1817.29	38 13.60	22 4.94	6.85	0.8	2.35	0.10	0.5	0.6	910827	3 1510.05	38 14.61	22 5.10	7.53	0.5	2.42	0.14	0.2	0.2	842
910825	20 3846.25	38 14.25	22 5.46	7.04	0.3	2.41	0.12	0.3	0.2	910827	3 5411.37	38 24.32	21 55.82	8.59	0.8	2.68	0.12	0.4	0.3	843
910825	23 10 6.98	38 24.75	22 8.13	10.99	0.5	2.39	0.04	0.5	0.4	910827	4 828.08	38 18.02	22 9.02	6.10	0.5	2.46	0.08	0.3	0.1	844
910826	0 1338.87	38 34.48	21 54.47	14.30	0.5	2.35	0.10	0.2	0.4	910827	4 5822.36	38 22.86	22 8.53	7.02	0.8	2.87	0.20	0.6	0.7	845
910826	0 16 7.49	38 22.14	21 53.48	6.68	0.7	2.24	0.09	0.3	0.3	910827	5 4953.42	38 23.91	21 49.22	9.85	1.7	2.81	0.05	0.9	0.6	846
910826	1 3214.75	38 24.74	21 58.40	8.79	1.9	2.27	0.12	1.7	0.5	910827	8 2640.22	38 14.67	22 5.89	8.40	0.8	2.46	0.09	0.3	0.3	847
910826	1 4317.41	38 24.27	21 57.77	7.13	0.4	1.94	0.09	0.2	0.2	910827	8 5250.32	38 18.83	22 4.72	8.98	0.8	2.65	0.09	0.5	0.2	848

APPENDIX C: FOCAL MECHANISMS

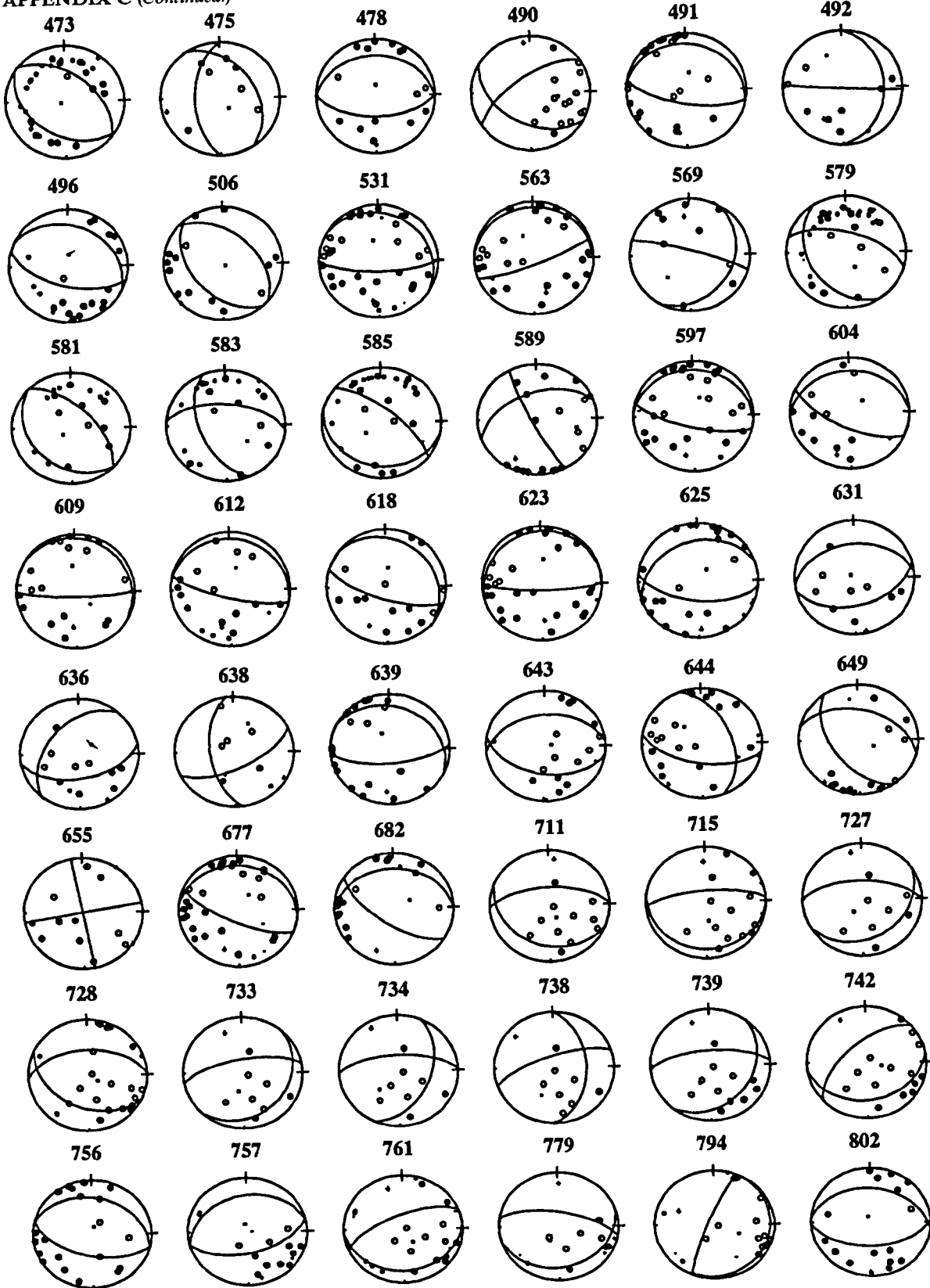
Lower hemisphere ●: compression - ○: dilatation - Smallest same dots for refracted rays



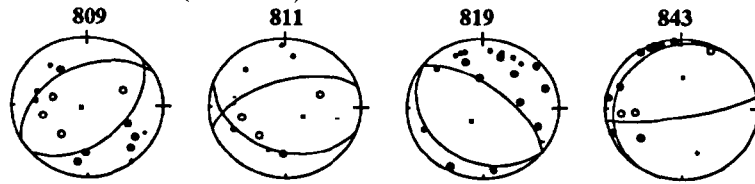
APPENDIX C (Continued.)



APPENDIX C (Continued.)



APPENDIX C (Continued.)



APPENDIX D: FOCAL MECHANISM PARAMETERS

Event	Str1	Dip1	Slip1	Str2	Dip2	Slip2
6	0.0	50.0	-139.9	241.6	60.5	-47.6
8	37.0	83.0	124.7	137.0	35.3	12.2
12	320.0	45.0	-112.2	170.0	49.1	-69.3
13	265.0	23.0	-105.7	102.0	67.9	-83.4
14	300.0	43.0	-129.1	168.0	58.0	-59.5
15	290.0	40.0	-70.0	84.6	52.8	-106.0
22	211.0	70.0	-150.4	110.0	62.3	-22.7
23	223.0	70.0	-157.7	125.0	69.1	-21.5
25	220.0	70.0	-168.4	126.0	79.2	-20.4
34	69.0	35.0	-90.0	249.0	55.0	-90.0
35	120.0	50.0	-90.0	300.0	40.0	-90.0
37	102.0	45.0	-88.6	280.0	45.0	-91.4
39	120.0	80.0	-19.9	213.6	70.4	-169.4
40	230.0	50.0	-99.8	65.0	41.0	-78.6
41	170.0	40.0	-97.9	0.3	50.5	-83.4
45	110.0	50.0	-79.6	274.0	41.1	-102.2
46	324.0	35.0	-45.7	94.0	65.8	-116.1
47	300.0	20.0	-59.6	88.0	72.8	-100.4
50	89.0	60.0	-56.4	216.0	43.8	-133.8
52	51.0	51.0	-93.2	236.0	39.1	-86.1
59	70.0	80.0	-110.0	314.5	22.3	-27.3
64	310.0	48.0	-34.9	65.0	64.9	-132.3
70	290.0	65.0	-94.3	120.0	25.3	-80.9
72	285.0	65.0	-87.9	100.0	25.1	-94.5
73	70.0	70.0	-90.0	250.0	20.0	-90.0
79	120.0	80.0	180.0	210.0	90.0	0.0
86	35.0	70.0	165.7	130.0	76.5	20.6
88	210.0	90.0	0.0	300.0	90.0	180.0
91	75.0	27.0	-90.0	255.0	63.0	-90.0
92	270.0	60.0	-120.0	139.1	41.4	-49.1
97	30.0	40.0	-60.0	173.0	56.2	-112.8
98	110.0	80.0	-90.0	290.0	10.0	-90.0
103	267.0	28.0	-99.7	98.0	62.4	-84.9
106	90.0	80.0	-100.0	315.4	14.1	-45.5
117	210.0	60.0	0.0	300.0	90.0	180.0
121	250.0	44.0	-78.3	54.0	47.1	-101.0
127	290.0	70.0	-100.0	137.3	22.3	-64.5
136	277.0	44.0	-56.1	54.0	54.8	-118.3
140	311.0	45.0	-102.2	148.0	46.3	-78.1
142	96.0	69.0	-85.6	264.0	21.4	-101.2
146	101.0	59.0	-53.7	226.0	46.3	-134.6
154	70.0	70.0	-80.1	223.0	22.2	-115.3
155	84.0	71.0	-96.0	282.0	19.9	-73.0
156	73.0	65.0	-118.4	305.0	37.1	-44.4
175	48.0	47.0	-98.2	240.0	43.6	-81.3
176	106.0	71.0	-100.6	316.0	21.7	-61.8

APPENDIX D (Continued.)

Event	Str1	Dip1	Slip1	Str2	Dip2	Slip2	Event	Str1	Dip1	Slip1	Str2	Dip2	Slip2
188	256.0	45.0	-104.4	96.0	46.8	-76.0	478	90.0	60.0	-90.0	270.0	30.0	-90.0
195	130.0	87.0	-66.9	227.0	23.3	-172.4	490	120.0	45.0	-34.6	236.0	66.3	-129.5
197	311.0	11.0	-46.5	87.0	82.0	-97.6	491	100.0	70.0	-90.0	280.0	20.0	-90.0
199	110.0	39.0	-49.2	242.0	61.5	-117.9	496	295.0	30.0	-81.3	105.0	60.4	-95.0
223	290.0	50.0	-79.6	94.0	41.1	-102.2	506	312.0	45.0	-95.7	140.0	45.3	-84.4
226	330.0	70.0	-100.0	177.3	22.3	-64.5	531	90.0	75.0	-99.2	302.0	17.5	-59.2
232	220.0	20.0	-150.0	101.5	80.1	-72.5	563	270.0	10.0	-70.0	69.7	80.6	-93.5
235	10.0	50.0	-140.0	251.7	60.5	-47.6	569	286.0	83.0	-62.2	29.0	28.6	-165.3
241	260.0	40.0	-149.8	146.0	71.2	-54.0	579	286.0	61.0	-113.6	148.0	36.7	-54.2
242	260.0	40.0	-150.0	146.1	71.2	-54.0	581	316.0	59.0	-89.0	134.0	31.0	-91.7
249	321.0	36.0	-70.2	117.0	56.4	-103.8	583	275.0	58.0	-136.1	158.0	54.0	-40.9
250	30.0	89.0	-149.0	299.4	59.0	-1.2	585	310.0	70.0	-80.1	103.0	22.2	-115.3
254	282.0	23.0	-54.3	64.0	71.5	-103.9	589	250.0	50.0	-169.9	153.5	82.3	-40.4
264	112.0	53.0	-90.0	292.0	37.0	-90.0	597	274.0	17.0	-99.6	104.0	73.2	-87.1
267	80.0	60.0	-120.0	309.1	41.4	-49.1	604	268.0	25.0	-115.7	116.0	67.6	-78.6
269	276.0	45.0	-92.8	100.0	45.1	-87.2	609	300.0	9.0	-60.3	90.0	82.2	-94.5
273	90.0	65.0	-90.0	270.0	25.0	-90.0	612	281.0	13.0	-92.9	104.0	77.0	-89.3
279	140.0	80.0	-59.9	246.7	31.5	-160.6	618	310.0	20.0	-69.2	108.0	71.4	-97.4
294	30.0	65.0	168.3	125.0	79.4	25.5	623	250.0	10.0	-110.0	90.3	80.6	-86.5
297	250.0	70.0	-169.9	156.5	80.5	-20.3	625	97.0	61.0	-76.7	251.0	31.7	-112.5
305	80.0	77.0	-104.0	308.0	19.0	-43.6	631	264.0	41.0	-83.9	76.0	49.3	-95.2
308	317.0	62.0	14.7	220.0	77.1	151.2	636	80.0	53.0	-63.2	220.0	44.5	-120.9
309	296.0	49.0	-100.7	132.0	42.1	-78.0	638	65.0	65.0	-37.6	173.0	56.5	-149.5
312	90.0	70.0	-100.0	297.3	22.3	-64.5	639	86.0	69.0	-101.7	296.0	23.9	-62.2
315	59.0	68.0	-146.4	315.0	59.1	-25.9	643	265.0	43.0	-101.1	100.0	48.0	-79.8
319	108.0	39.0	-75.8	270.0	52.4	-101.2	644	100.0	60.0	-130.0	339.2	48.4	-41.9
322	300.0	20.0	-60.0	88.4	72.8	-100.3	649	285.0	45.0	-121.6	146.0	53.0	-62.4
324	315.0	48.0	-111.9	166.0	46.4	-67.5	655	80.0	90.0	180.0	170.0	90.0	0.0
326	300.0	21.0	-71.2	100.0	70.2	-97.0	677	276.0	19.0	-107.1	114.0	71.9	-84.2
335	256.0	47.0	-73.1	52.0	45.6	-107.3	682	269.0	29.0	-121.5	124.0	65.6	-73.9
337	332.0	69.0	-21.4	70.0	70.1	-157.6	711	112.0	29.0	-68.7	268.0	63.1	-101.4
349	240.0	60.0	-148.6	133.0	63.1	-34.1	715	94.0	17.0	-82.3	266.0	73.2	-92.3
350	90.0	80.0	-90.0	270.0	10.0	-90.0	727	58.0	31.0	-118.2	270.0	63.0	-74.2
356	330.0	66.0	-9.8	64.0	81.1	-155.7	728	116.0	41.0	-68.1	268.0	52.5	-107.9
358	350.0	45.0	46.8	223.0	59.0	124.4	733	42.0	25.0	-125.3	260.0	69.8	-74.9
360	75.0	73.0	-128.8	325.0	41.8	-26.0	734	24.0	45.0	-145.4	268.0	66.3	-50.5
364	36.0	85.0	-158.2	304.0	68.3	-5.4	738	0.0	45.0	-155.3	252.0	72.8	-47.7
373	308.0	45.0	-87.2	124.0	45.1	-92.8	739	38.0	31.0	-133.6	266.0	68.1	-67.5
377	300.0	30.0	-10.0	38.7	85.0	-119.6	742	80.0	33.0	-63.3	229.0	60.9	-106.3
385	85.0	68.0	-144.4	340.0	57.4	-26.4	756	284.0	33.0	-90.0	104.0	57.0	-90.0
398	115.0	73.0	-72.0	247.0	24.6	-135.3	757	76.0	37.0	-90.0	256.0	33.0	-90.0
403	80.0	75.0	-145.7	340.0	57.1	-18.0	761	85.0	23.0	-74.3	248.0	67.9	-96.6
409	95.0	50.0	-73.3	250.0	42.8	-108.9	779	72.0	25.0	-115.7	280.0	67.6	-78.6
411	81.0	59.0	-98.9	278.0	32.1	-75.5	794	28.0	9.0	-85.1	203.0	81.0	-90.8
422	80.0	25.0	-90.0	260.0	65.0	-90.0	802	276.0	37.0	-83.6	88.0	53.3	-94.8
423	55.0	45.0	-67.8	205.0	49.1	-110.7	809	47.0	47.0	-98.9	240.0	43.7	-80.5
431	310.0	30.0	-60.0	96.3	64.3	-106.1	811	110.0	37.0	-58.0	252.0	59.3	-111.7
437	84.0	65.0	-75.2	232.0	28.8	-118.7	819	308.0	65.0	-90.0	128.0	25.0	-90.0
447	350.0	30.0	-40.0	116.0	71.3	-113.9	843	80.0	80.0	-80.1	215.0	14.0	-134.1
463	45.0	68.0	-2.7	136.0	87.5	-158.0							
473	124.0	40.0	-93.8	309.0	50.1	-86.8							
475	180.0	55.0	-69.6	327.0	39.9	-116.5							